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# ADVANCES IN IONOSPHERIC MAPPING BY NUMERICAL METHODS

WILLIAM B. JONES, RONALD P. GRAHAM AND MARGO LEFTIN



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U. S. DEPARTMENT OF COMMERCE  
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# NATIONAL BUREAU OF STANDARDS

## *Technical Note 337*

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### ADVANCES IN IONOSPHERIC MAPPING BY NUMERICAL METHODS

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# ADVANCES IN IONOSPHERIC MAPPING BY NUMERICAL METHODS

by

William B. Jones, Ronald P. Graham and Margo Leftin

Paper describes recent progress made at ITSA (formerly CRPL) to improve numerical methods for mapping and predicting ionospheric characteristics used in telecommunication. Two major problems are considered: (1) tendency of maps to smooth out physical properties of the ionosphere, particularly at low latitudes and (2) ambiguous values at geographic poles and resulting distortions in immediate surroundings. The second problem is overcome by means of a universal time analysis. Significant improvement is made in the first problem by the use of a modified magnetic dip coordinate. A general description is given of the new procedures for forming numerical maps, including a number of illustrations. Also included are a discussion of the use of numerical maps in the revised form and a Fortran Program for evaluating numerical maps of ionospheric characteristics and for calculating the  $MUF(ZERO)F2$  and  $MUF(4000)F2$ .

## 1. Introduction

In recent years the Institute for Telecommunication Sciences and Aeronomy (formerly the Central Radio Propagation Laboratory) has been engaged in the development of numerical methods and computer programs for mapping and predicting characteristics of the ionosphere used in telecommunication. Work along these lines was initiated prior to the International Geophysical Year, with the first descriptive publication appearing in the Telecommunication Journal [Jones and Gallet, 1960]. Since that time numerous technical papers have been written to describe the further advances in: (a) methods for data analysis and representation of ionospheric characteristics [Jones, 1962; Jones and Gallet, 1962a, 1965a, 1965b], (b) methods for applying numerical maps of ionospheric characteristics [Jones and Gallet, 1962b], and

(c) development of computer programs associated with this work [Hinds and Jones, 1963]. The methods described in the preceding references have been used as the basis for a number of applications in ionospheric predictions: see for example [Haydon and Lucas, in publication; Lucas and Haydon, in publication; Davies, 1965]. In January 1963 the methods were adopted for producing Ionospheric Predictions [ITSA], issued monthly by the Institute for Telecommunication Sciences and Aeronomy (formerly CRPL) as an aid to determining the best sky wave frequencies over any transmission path, at any time of day, for average monthly conditions.

Although the previous work on ionospheric mapping (referred to above) has overcome many of the basic mapping problems and produced numerical representations and predictions of ionospheric characteristics which are useful and acceptable for many purposes, it was recognized [Jones and Gallet, 1962a, 1962b, 1965a] that there remained certain weaknesses in numerical methods of mapping which would eventually need further work. The present paper is concerned with recent efforts focused on two such problem areas: (1) a tendency of the maps to smooth out certain physical properties of the ionosphere, particularly at low latitudes, and (2) the presence of ambiguous values in the maps at geographic poles, resulting from a local mean time analysis. The second problem has been overcome by means of an analysis in universal time. Evidence of the first problem is found in residuals (between actual data and computed values) which are above the level of noise inherent in ionospheric measurements. Significant progress in this



area has been achieved by taking into account more explicitly the relationship between ionospheric characteristics and the magnetic field. Previous studies<sup>1</sup> have indicated the dependence of the ionosphere on the magnetic field, particularly at low latitudes.

The revised methods and programs described in this paper have been used to produce a provisional Atlas of Ionospheric Characteristics for consideration at the 12th Plenary Assembly of CCIR to be held at Oslo, Norway, June 1966. For this Atlas, two characteristics have been analyzed: monthly median of the F2 layer critical frequency ( $f_oF_2$ ) for all months of the five year period extending from 1954 through 1958, and monthly median of  $M(3000)F_2$  for all months of 1954 and 1958. The revised method will also soon be used as the basis for ITSA Ionospheric Predictions.

The intention of the present paper is to give a general description of the new procedures for ionospheric mapping, with emphasis on illustrating various main steps by means of graphs, maps, and tables. Also included are: (1) a comparison of results from two different types of analyses, (2) a careful discussion of how to apply numerical maps of ionospheric characteristics in their present (revised) form and (3) a FORTRAN program for evaluating numerical maps of ionospheric characteristics and for calculating  $MUF(ZERO)F_2$  and  $MUF(4000)F_2$  (Appendix D). By way of illustration we have used only the characteristic,  $f_oF_2$  median, since it is the most interesting and important for telecommunication.

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<sup>1</sup>[GHQ United States Army, 1946; Appleton, 1946; Bailey, 1948]

## 2. Screen Analysis

As in the previous paper by Jones and Gallet [1965a] the numerical mapping procedures described here involve two principal analyses of the data, both performed in one pass on the computer. First, a so-called "screen analysis" is made of available A- and B-data<sup>1</sup> to produce approximate values of the ionospheric characteristic (called "C-data") at carefully chosen locations ("screen points") in large regions, such as oceans, where no stations are available. These regions will be called the critical areas. The C-data are then combined with the original A- and B-data in forming a second (and final) analysis. The purpose of C-data, discussed at length in the preceding reference, is to control instabilities in the maps, which tend to develop in the critical areas.

The screen analysis is started by Fourier analyzing the monthly medians of foF2 for each station for the 24 hours of the day. Thus the diurnal variation at each place is decomposed into eleven harmonics

$$a_j \cos jt + b_j \sin jt = c_j \cos(jt - \psi_j), \quad 1 \leq j \leq 11, \quad (1)$$

and a constant term  $a_0$  which is the arithmetic mean of the 24 medians<sup>2</sup>.

Here  $t$  denotes the local mean hour angle ( $-180^\circ \leq t \leq 180^\circ$ ,  $t = 0^\circ$  at noon LMT). Each harmonic is defined by an amplitude  $c_j$  and phase  $\psi_j$ ,

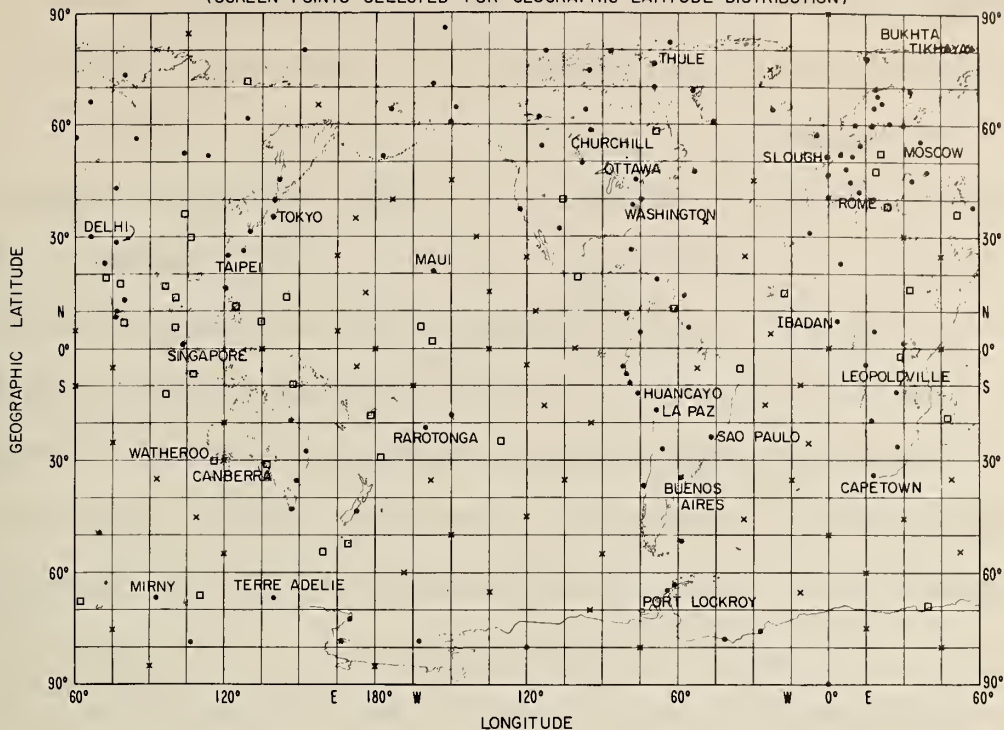
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<sup>1</sup>The terms A- and B-data are defined in the above references. In short, A-data refer to measurements at stations taken during the actual month in question. B-data refer to values of the characteristic obtained from interpolation or extrapolation in time at a station which did not report data for the specific month in question but was active for a period of time before or after. B-data are used to fill gaps where A-data are not available (Fig. 1).

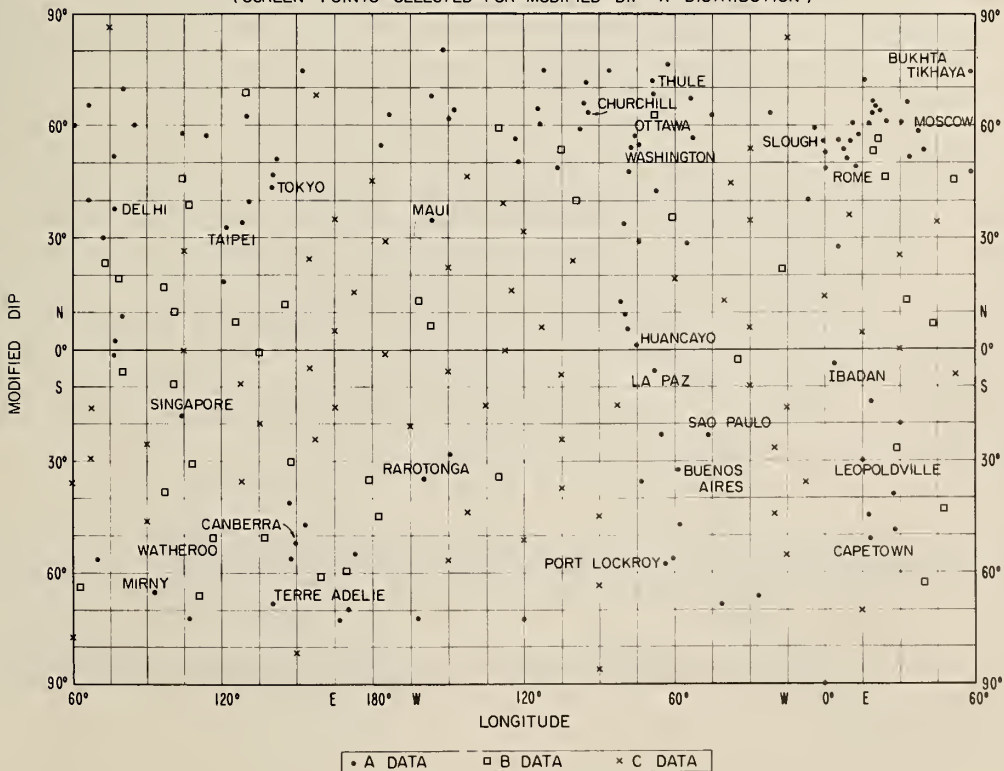
<sup>2</sup>There is also an additional term  $a_{12} \cos 12t$  which is omitted since it does not form a complete harmonic and is produced mainly by noise.

Figure 1

MAP OF IONOSPHERIC STATIONS (MARCH 1958)  
(SCREEN POINTS SELECTED FOR GEOGRAPHIC LATITUDE DISTRIBUTION)



MAP OF IONOSPHERIC STATIONS (MARCH 1958)  
(SCREEN POINTS SELECTED FOR MODIFIED DIP X DISTRIBUTION)



• A DATA    □ B DATA    × C DATA

equivalently, by cosine and sine coefficients  $a_j$  and  $b_j$ , respectively. Harmonics above the eighth order are dropped since they are produced more by noise than by real physical variation [Jones, 1962; Jones and Gallet, 1962a and 1965b ].

Basically, the purpose of the screen analysis is to reduce the major part of the two-dimensional geographic variation of each Fourier coefficient  $a_j$  and  $b_j$  ( $j \leq 8$ ) to that of a single variable  $x$ , which in general may be a function of both geographic latitude  $\lambda$  and longitude  $\theta$ . Since there will be no large gaps in the distribution of stations in the one dimension of  $x$ , an approximate representation of the ionospheric characteristic can be obtained as a function of  $x$  without the danger of instability. Accordingly, the Fourier coefficients  $a_j$  and  $b_j$  are represented geographically (to a first approximation) by a series of the form

$$\sum_{k=0}^K D_k G_k(x), \quad (2)$$

where the functions  $G_k(x)$  may be powers of  $x$  or  $\sin x$ , or some other functions. The coefficients  $D_k$  are determined by the method of least squares and the cutoff  $K$  is made (by the computer) using a Student's  $t$  test [Jones and Gallet, 1962a]. Thus the approximate representation of the ionospheric characteristic has the form

$$\Gamma(x, t) = a_0(x) + \sum_{j=1}^8 \left[ a_j(x) \cos jt + b_j(x) \sin jt \right], \quad (3)$$

where  $a_j(x)$  and  $b_j(x)$  represent the Fourier coefficients as functions of  $x$  in the form (2).

In previous work [Jones and Gallet, 1965a] the variable  $x$  was taken to be geographic latitude  $\lambda$ , and  $\sin^k \lambda$  was selected for the function  $G_k(x)$ . This choice of  $G_k$  was justified by a study included in Jones and Gallet, [1962a]. Results from this method were reasonably good, since the Fourier coefficients  $a_j$  and  $b_j$  are adjusted to local mean time and hence their main geographic variation is latitudinal. The resulting errors in C-data were therefore due to neglected secondary effects of longitude and, to a lesser extent, noise inherent in the measurements. A quantity used to measure the error in C-data is the root mean square of residuals between A- and B-data and values computed by (3). Values of this quantity were given for a representative sample of months by Jones and Gallet [1965a].

One of the main aspects of the present study has centered around the question of finding a best (or at least better) choice of  $x$ . Most of the work has been concentrated on variables related to the magnetic field, because of its strong effect on ionospheric characteristics. With geomagnetic latitude for the variable  $x$ , no noticeable improvement was achieved. On the other hand, by taking  $x$  to be the magnetic dip  $I$ , and certain simple modifications of  $I$ , significant reductions in residuals of the screen analysis have resulted<sup>1</sup>. This will be illustrated by comparing screen analyses of the same data using for  $x$  the five variables:  $\lambda$ ,  $I$ , and the following three modifications of  $I$  suggested by various workers<sup>2</sup>:

---

<sup>1</sup>The functional dependence of  $I$  on  $\lambda$  and  $\theta$  will be defined in Section 5, using results from the Jensen and Cain [1962] analysis of the magnetic field for epoch 1960.

<sup>2</sup>The basic parameter  $\frac{I}{\sqrt{\cos \lambda}}$  was suggested by K. Rawer.

$$x_1 = \text{Tan}^{-1} \left[ \frac{I}{\sqrt{\cos \lambda}} \right], \quad (4)$$

$$x_2 = \text{Tan}^{-1} \left[ \frac{\tan I}{\sqrt{\cos \lambda}} \right], \quad (5)$$

$$x_3 = \text{Tan}^{-1} \left[ \frac{\tan I}{\sqrt{\cos (\lambda - \delta)}} \right], \quad (6)$$

where  $\delta$  denotes solar declination<sup>1</sup>. In each analysis  $G_k(x) = \sin^k x$  was used with a maximum possible value of 12 for the degree  $K$ .

The root mean square (RMS) of residuals between foF2 medians (A- and B-data) and values computed from (3) for all hours and all stations<sup>2</sup> are given in Table 1, for four seasonal months of the high and low years of solar activity 1958 and 1954. From these results may be seen a strong improvement with the use of  $I$  compared with  $\lambda$ , for all months except June, at which time the difference is insignificant. In addition, the modified variables  $x_1, x_2, x_3$  give some improvement over  $I$ , particularly during high solar activity (1958). Although the difference between values of the RMS residual obtained using  $x_1, x_2$  and  $x_3$  is small, the values for  $x_1$  appear slightly smaller in most months. Hence  $x_1$  was selected as the best geographic variable for the screen analysis.

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<sup>1</sup>Here, as throughout the text, we use  $\text{Tan}_0^{-1}$  to denote the principal value  $v$  of the arctangent such that  $-90^\circ \leq v \leq 90^\circ$ .

<sup>2</sup>The approximate numbers of stations with A- and B-data for each month are given in Table 6.

Table 1  
 RMS Residuals (Mc/s) from Screen Analyses  
 of foF2 Medians

Variable (x)	1954				1958			
	Mar.	June	Sept.	Dec.	Mar.	June	Sept.	Dec.
$\lambda$	.85	.68	.80	.78	1.41	1.25	1.35	1.31
I	.65	.69	.68	.69	1.06	1.24	1.00	1.08
$x_1$	.62	.63	.65	.67	.97	1.08	.95	1.02
$x_2$	.61	.65	.65	.68	.98	1.12	.95	1.05
$x_3$	.62	.64	.65	.70	.98	1.14	.95	1.19

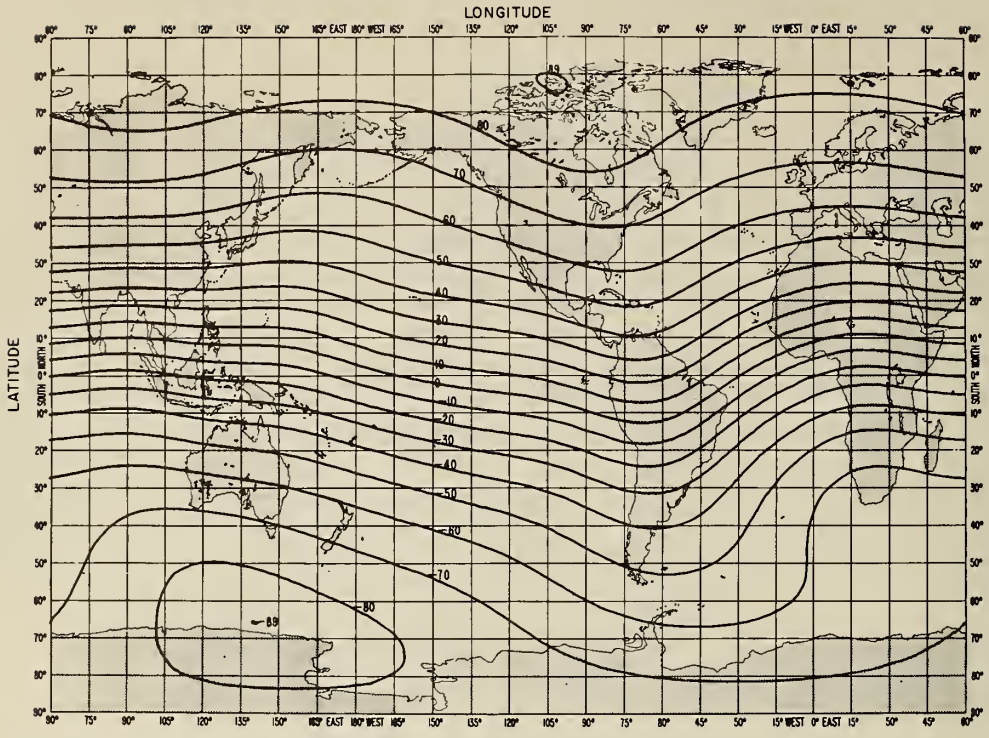
( $\lambda$  = geographic latitude, I = magnetic dip,  
 $x_1$ ,  $x_2$ , and  $x_3$  are modified forms of I.)

In the sequel  $x_1$  will be the only modified form of I used; hence from this point on we shall drop the subscript and write  $x_1$  as merely x. To help understand the difference between x and I, contour maps of these two quantities are shown in figure 2. The main difference occurs in the higher latitudes. At low latitudes x is quite similar to I, but at high latitudes it tends to behave more as geographic latitude  $\lambda$ . Thus, in a sense, x combines certain characteristics of both I and  $\lambda$ . Perhaps this accounts for the superiority of x over  $\lambda$  and I as the geographic variable.

A study of the distribution of Fourier coefficients  $a_j$  and  $b_j$ , relative to the various geographic variables considered above, throws much light on the significance of Table 1 and helps explain the advantages of certain choices of the geographic variable over others. As an example,

Figure 2

MAGNETIC DIP  
EPOCH 1960 HEIGHT 300 km



MODIFIED MAGNETIC DIP  
EPOCH 1960 HEIGHT 300 km

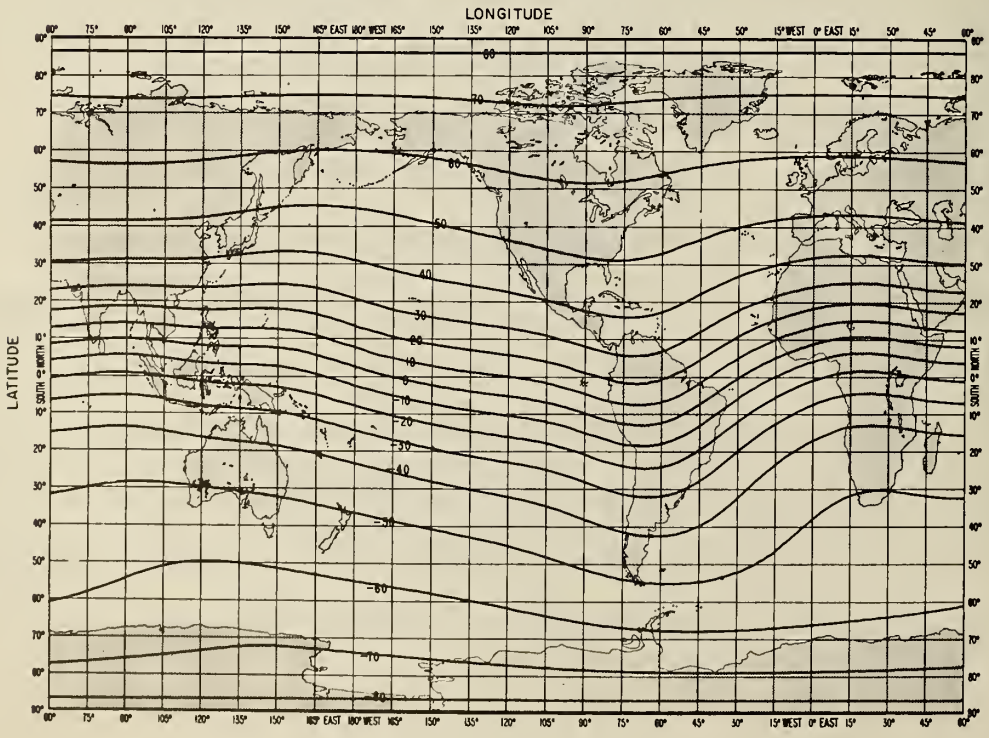




Figure 3 shows the distribution of  $a_0$  (the arithmetic mean of the 24 hour medians of foF2) as a function of geographic latitude (top) and of the modified magnetic dip  $x$  (bottom) for March 1958. Here, as expected from previous studies, we find a remarkable simplification induced by the variable  $x$ . On each graph is shown the least square approximation of the form (2) discussed above. The graphs are made somewhat three-dimensional by the use of special plotting symbols to distinguish the approximate geographic longitudes of stations (see Longitudinal Code on each graph). From this it can be seen that the Fourier coefficients are not randomly dispersed about the least square curve; in certain intervals they are systematically grouped according to longitude. Similar illustrations for the coefficients  $a_j$  and  $b_j$ ,  $j = 1, 2, 3, 4$ , are contained in Appendix A. For the first two harmonics the advantage of  $x$  over  $\lambda$  is seen to be comparable with that in the case of  $a_0$ . For higher harmonics the difference appears to be less noticeable.

From the standpoint of the screen analysis, the most important aspect of the distributions illustrated is the dispersion of the coefficients about the least square curves. Quantitatively, this dispersion is measured in terms of standard deviation of the residuals, an unbiased estimate of which is computed by

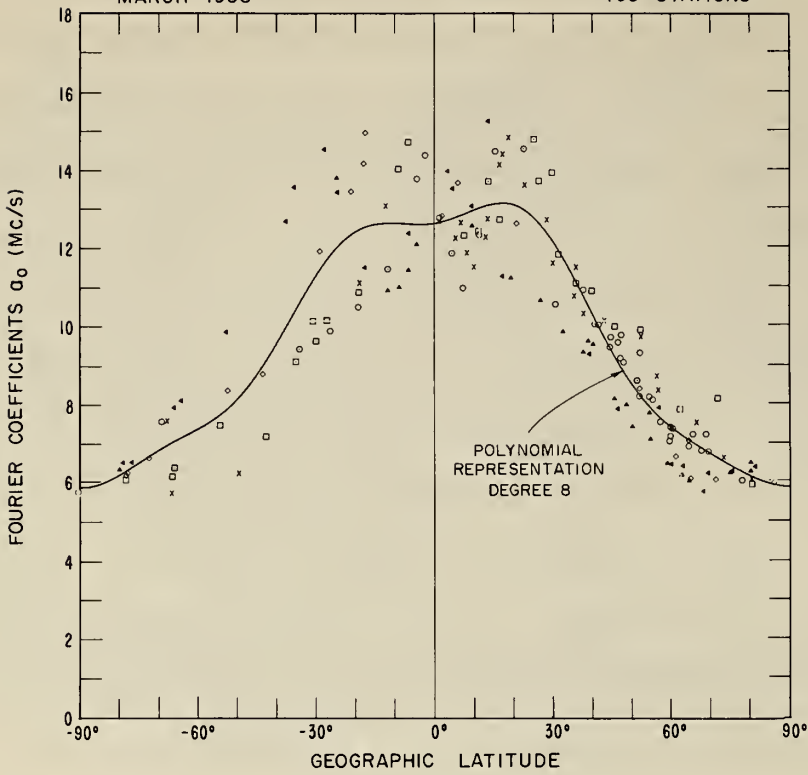
$$e_K = \left[ \frac{E_K}{N - K - 1} \right]^{\frac{1}{2}} \quad (7)$$

where  $E_K$  denotes the sum of squares of the residuals,  $N$  is the number of stations with A- or B-data, and  $(N - K - 1)$  is the number of degrees of freedom remaining after subtraction of 1 for each function in the series

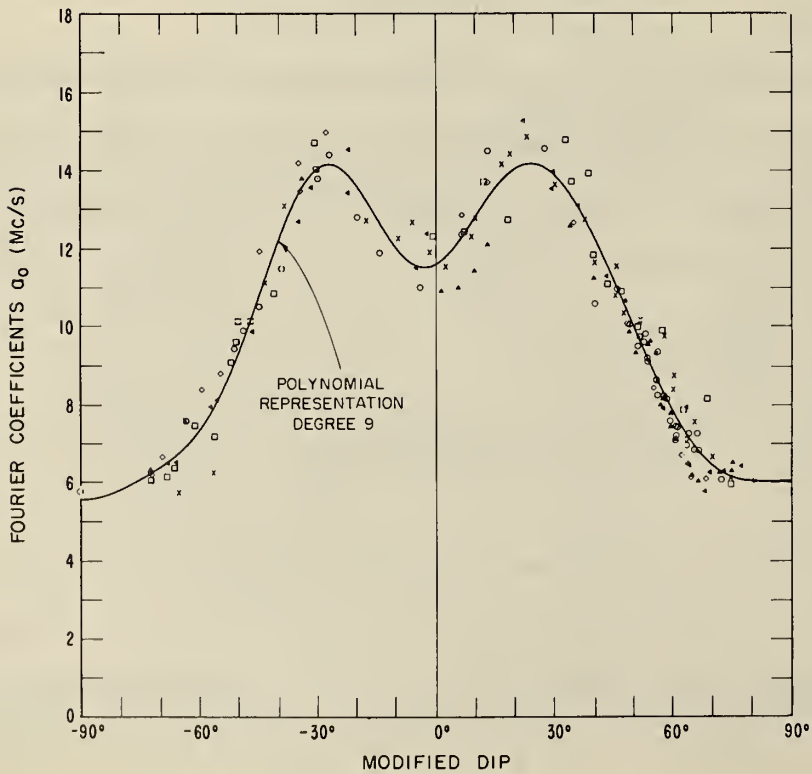
Figure 3

FOURIER COEFFICIENTS FOR foF2 MONTHLY MEDIANS  
MARCH 1958 163 STATIONS

$a_0$



LONGITUDINAL CODE	
○	345°E - 45°E
x	45°E - 105°E
□	105°E - 165°E
◇	165°E - 225°E
△	225°E - 285°E
◻	285°E - 345°E



(2). Values of  $e_K$  are given in Table 2 for each Fourier coefficient through harmonic 4, March 1958, from analyses with  $\lambda$  and with  $x$ . The results shown in Table 2 are seen to agree with evidence brought out in the graphical illustrations: there is a great improvement from the use of  $x$  which lies mainly in the first two harmonics and the constant term  $a_0$ .

Table 2  
Standard Deviation of Residuals  $e_K$  (Mc/s) from  
Analyses of Fourier Coefficients for March 1958.

Variable	$a_0$	$a_1$	$b_1$	$a_2$	$b_2$	$a_3$	$b_3$	$a_4$	$b_4$
$\lambda$	1.09	.75	.74	.42	.43	.23	.23	.13	.13
$x$	.64	.54	.54	.35	.36	.20	.18	.13	.12

( $\lambda$  = geographic latitude,  $x$  = modified magnetic dip)

As a final illustration of the screen analysis, figures 4 and 5 show maps of foF2 median for March 1958 computed from screen analysis representations of the form (3), using  $\lambda$  (top) and  $x$  (bottom) for the two hours U.T. = 06 and 18. As expected, the lower maps (from analysis with  $x$ ) give evidence of a strong connection with the magnetic field, whereas the upper maps (from analysis with  $\lambda$ ) appear more symmetric with respect to the geographic equator. The lower maps represent a much greater degree of detail in the equatorial region. Attention is called to a further point illustrated by these maps: the presence of ambiguous values at the geographic poles. As an example, the lower map for U.T. = 18 has an

Figure 4

MAPS OF foF2 MEDIAN (Mc/s) FROM SCREEN ANALYSIS

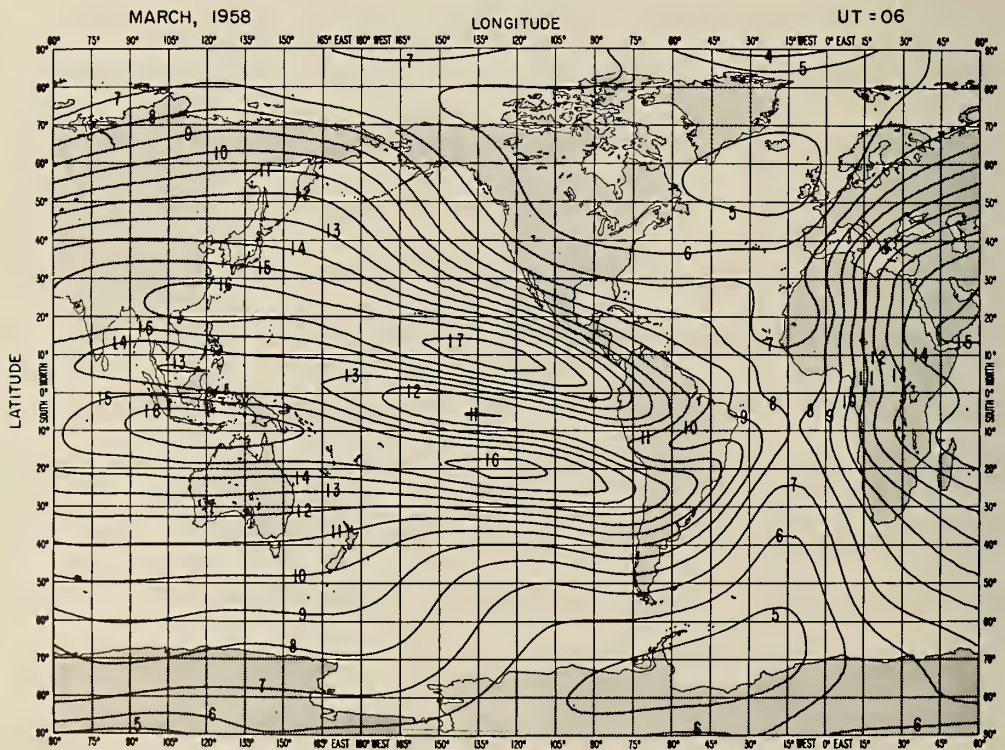
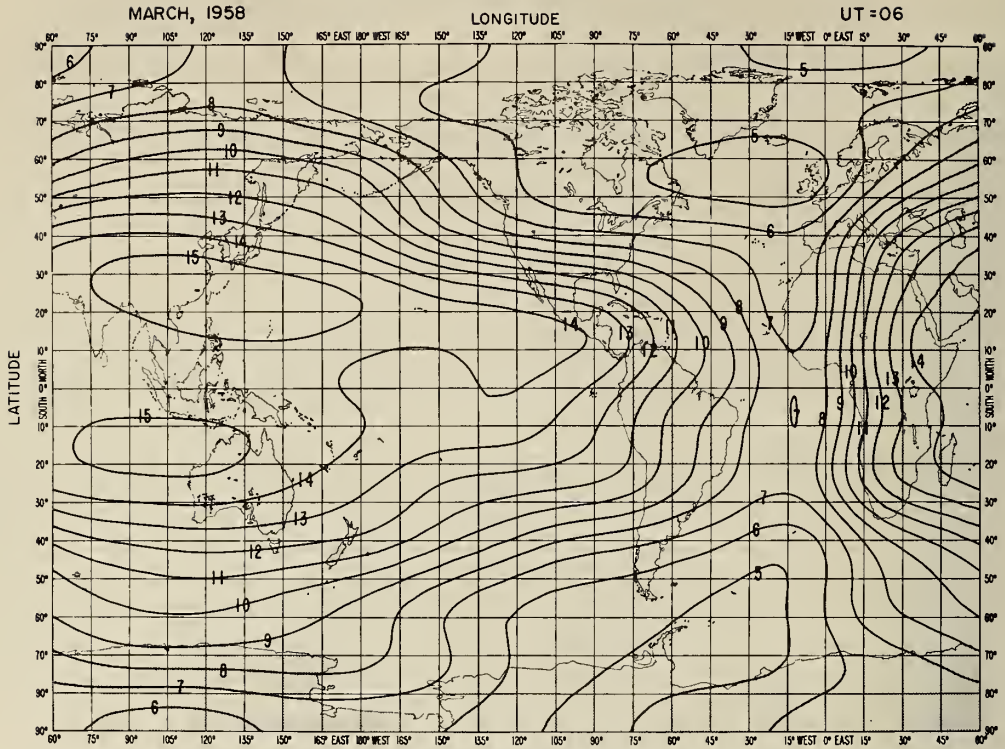
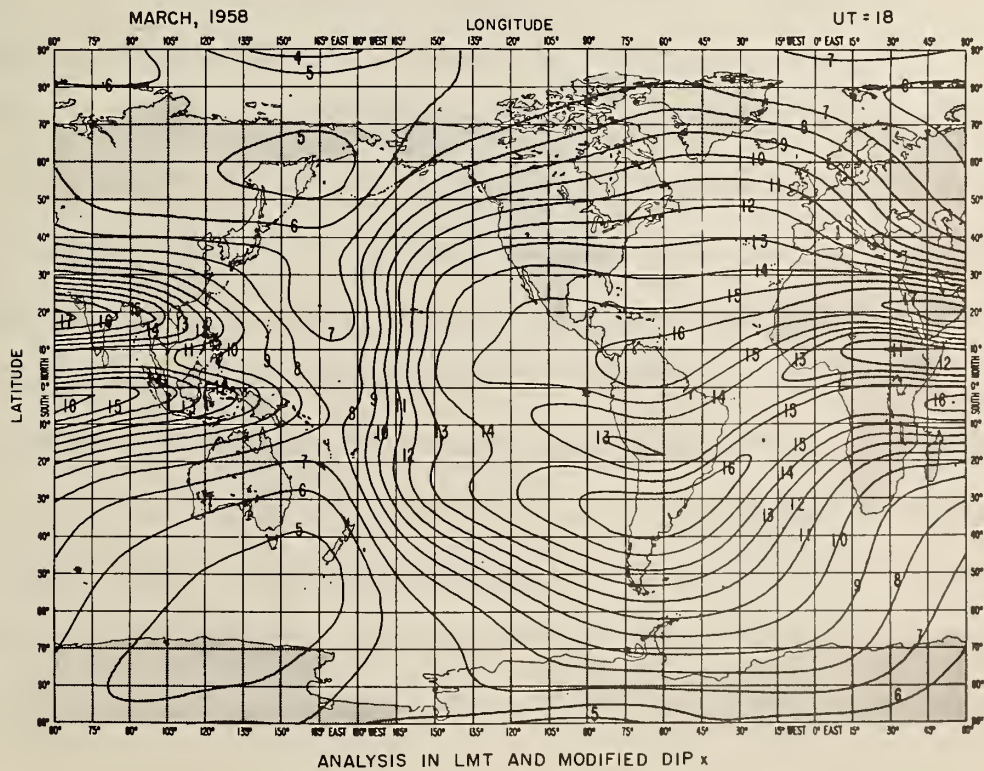
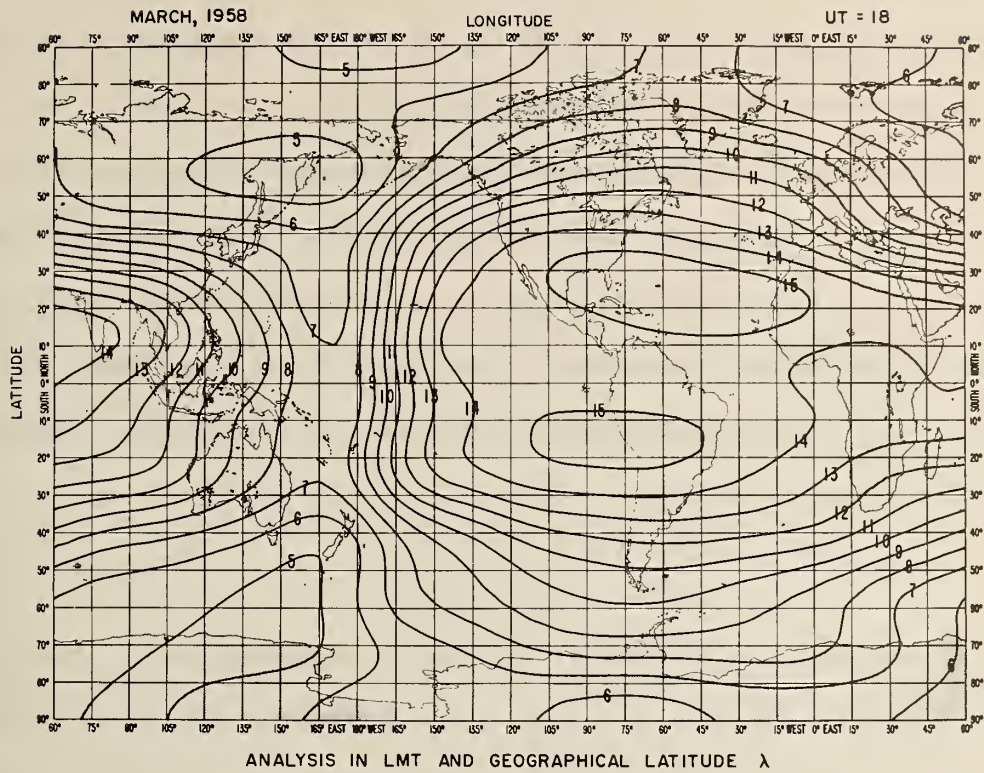


Figure 5

MAPS OF foF2 MEDIAN (Mc/s) FROM SCREEN ANALYSIS



interval of values from less than 4 to greater than 7 Mc/s at the North Pole. As mentioned in the introduction, the polar problem is a consequence of the analysis made in local mean time. It affects not only the poles but also produces distortion in small neighborhoods about them. The problem is not of great importance in the screen analysis, since the representation from the screen analysis is used only for defining C-data at screen points which are rarely needed in the close vicinity of the poles. The nature of the polar representation is important in the second analysis, however, and for that we have adopted the approach described in the following section.

### 3. Second Analysis

The second analysis is made of the combined set of A-, B-, and C-data. As a result, improvements in the screen analysis are carried over to the numerical map produced by the second analysis. The chief concern of this section is with basic differences in procedures for the second analysis described here and that described by Jones and Gallet [1965a]. The most important are the following: (1) reversal of the order of diurnal and geographic analyses, (2) geographic analyses of A-, B-, and C-data for fixed instants of universal time, (3) change in geographic functions to include explicit dependence on the modified magnetic dip  $x$ , as well as on  $\lambda$  and  $\theta$ , and (4) diurnal analysis of orthonormal coefficients derived from the geographic analyses.

The use of modified magnetic dip  $x$  as a principal coordinate has resulted in a change in selecting screen points. As mentioned earlier, screen points are chosen to fill wide gaps in the distribution of stations

where instabilities might develop. The use of  $x$  and  $\theta$  as principal coordinates (instead of  $\lambda$  and  $\theta$ ) has the effect of transforming the station distribution and hence of changing the shape and size of the gaps. Thus it is necessary to select different screen points from those used in previous analyses. As an illustration, the station distribution for March 1958 is shown in figure 1(a) in  $(\lambda, \theta)$  coordinates and figure 1(b) in  $(x, \theta)$  coordinates. The sets of screen points selected for the former distribution and for the transformed distributions are shown in the respective maps by the use of special plotting symbols (see captions). As in the previous case, the set of screen points changes very little from month to month.

### 3.1 Geographic Variation

The world-wide geographic variations of monthly median foF2 for each of 24 fixed hours of universal time (U.T. = 00, 01, ..., 23), are represented by series of the form

$$\sum_{k=0}^{75} D_k G_k(\lambda, \theta), \quad (8)$$

where the geographic functions  $G_k(\lambda, \theta)$  are defined in Table 3. The coefficients  $D_k$  are determined by the method of least squares using all available A-, B-, and C-data for the given hour. The modified magnetic dip  $x$  does not appear in the functional form  $G_k(\lambda, \theta)$ , since  $x$  is itself a function of  $\lambda$  and  $\theta$  (see Section 5). The variable  $\lambda$  has been retained in the weighting factors involving powers of  $\cos \lambda$ , in order to provide a continuous latitudinal variation of the weight factors and to annihilate

Table 3

Geographic Coordinate Functions  $G_k(\lambda, \theta)$  for Analysis of foF2 Median  
 ( $\lambda$  = geographic latitude,  $\theta$  = geographic longitude,  $x$  = modified magnetic dip)

k	$G_k(\lambda, \theta)$	k	$G_k(\lambda, \theta)$
0	1	40	$\sin^2 x \cos^2 \lambda \cos 2\theta$
1	$\sin x$	41	$\sin^2 x \cos^2 \lambda \sin 2\theta$
2	$\sin^2 x$	42	$\sin^3 x \cos^2 \lambda \cos 2\theta$
3	$\sin^3 x$	43	$\sin^3 x \cos^2 \lambda \sin 2\theta$
4	$\sin^4 x$	44	$\sin^4 x \cos^2 \lambda \cos 2\theta$
5	$\sin^5 x$	45	$\sin^4 x \cos^2 \lambda \sin 2\theta$
6	$\sin^6 x$	46	$\sin^5 x \cos^2 \lambda \cos 2\theta$
7	$\sin^7 x$	47	$\sin^5 x \cos^2 \lambda \sin 2\theta$
8	$\sin^8 x$	48	$\sin^6 x \cos^2 \lambda \cos 2\theta$
9	$\sin^9 x$	49	$\sin^6 x \cos^2 \lambda \sin 2\theta$
10	$\sin^{10} x$	50	$\sin^7 x \cos^2 \lambda \cos 2\theta$
11	$\sin^{11} x$	51	$\sin^7 x \cos^2 \lambda \sin 2\theta$
		52	$\sin^8 x \cos^2 \lambda \cos 2\theta$
		53	$\sin^8 x \cos^2 \lambda \sin 2\theta$
12	$\cos \lambda \cos \theta$	54	$\cos^3 \lambda \cos 3\theta$
13	$\cos \lambda \sin \theta$	55	$\cos^3 \lambda \sin 3\theta$
14	$\sin x \cos \lambda \cos \theta$	56	$\sin x \cos^3 \lambda \cos 3\theta$
15	$\sin x \cos \lambda \sin \theta$	57	$\sin x \cos^3 \lambda \sin 3\theta$
16	$\sin^2 x \cos \lambda \cos \theta$	58	$\sin^2 x \cos^3 \lambda \cos 3\theta$
17	$\sin^2 x \cos \lambda \sin \theta$	59	$\sin^2 x \cos^3 \lambda \sin 3\theta$
18	$\sin^3 x \cos \lambda \cos \theta$	60	$\sin^3 x \cos^3 \lambda \cos 3\theta$
19	$\sin^3 x \cos \lambda \sin \theta$	61	$\sin^3 x \cos^3 \lambda \sin 3\theta$
20	$\sin^4 x \cos \lambda \cos \theta$	62	$\sin^4 x \cos^3 \lambda \cos 3\theta$
21	$\sin^4 x \cos \lambda \sin \theta$	63	$\sin^4 x \cos^3 \lambda \sin 3\theta$
22	$\sin^5 x \cos \lambda \cos \theta$	64	$\cos^4 \lambda \cos 4\theta$
23	$\sin^5 x \cos \lambda \sin \theta$	65	$\cos^4 \lambda \sin 4\theta$
24	$\sin^6 x \cos \lambda \cos \theta$	66	$\sin x \cos^4 \lambda \cos 4\theta$
25	$\sin^6 x \cos \lambda \sin \theta$	67	$\sin x \cos^4 \lambda \sin 4\theta$
26	$\sin^7 x \cos \lambda \cos \theta$	68	$\cos^5 \lambda \cos 5\theta$
27	$\sin^7 x \cos \lambda \sin \theta$	69	$\cos^5 \lambda \sin 5\theta$
28	$\sin^8 x \cos \lambda \cos \theta$	70	$\cos^6 \lambda \cos 6\theta$
29	$\sin^8 x \cos \lambda \sin \theta$	71	$\cos^6 \lambda \sin 6\theta$
30	$\sin^9 x \cos \lambda \cos \theta$	72	$\cos^7 \lambda \cos 7\theta$
31	$\sin^9 x \cos \lambda \sin \theta$	73	$\cos^7 \lambda \sin 7\theta$
32	$\sin^{10} x \cos \lambda \cos \theta$	74	$\cos^8 \lambda \cos 8\theta$
33	$\sin^{10} x \cos \lambda \sin \theta$	75	$\cos^8 \lambda \sin 8\theta$
34	$\sin^{11} x \cos \lambda \cos \theta$		
35	$\sin^{11} x \cos \lambda \sin \theta$		
36	$\cos^2 \lambda \cos 2\theta$		
37	$\cos^2 \lambda \sin 2\theta$		
38	$\sin x \cos^2 \lambda \cos 2\theta$		
39	$\sin x \cos^2 \lambda \sin 2\theta$		



functions involving  $\theta$ , at the geographic poles. This prevents further ambiguous values at the poles. The variable  $x$  could also have been used in the weight factors, in place of  $\lambda$ , since  $\cos x = 0$ , when  $\lambda = \pm 90^\circ$ . This change could be considered in future work, although its effect would perhaps be negligible.

In determining the least square approximations (8), a related set of orthonormal functions  $F_k(\lambda, \theta)$  is formed as described by Jones and Gallet [1962a]. These functions are linear combinations of the  $G_k(\lambda, \theta)$

$$F_k(\lambda, \theta) = \sum_{j=1}^{k-1} d_{kj} F_j(\lambda, \theta) + d_{kk} G_k(\lambda, \theta), \quad (9)$$

constructed so as to satisfy orthonormality conditions

$$\sum_{i=1}^N F_k(\lambda_i, \theta_i) F_j(\lambda_i, \theta_i) = \begin{cases} 0 & k \neq j \\ 1 & k = j \end{cases} \quad (10)$$

relative to the set of  $N$  positions  $(\lambda_i, \theta_i)$  of ionospheric stations and screen points. Thus the least square approximations (8) are obtained by series of the form

$$\sum_{k=0}^{75} d_k F_k(\lambda, \theta), \quad (11)$$

in which the terms are mutually independent and can be tested individually for statistical significance. By an algebraic transformation, (11) can be converted to the form (8). The procedures for this transformation are described in the above reference, but in the present application our interest will be in the coefficient  $d_k$ .

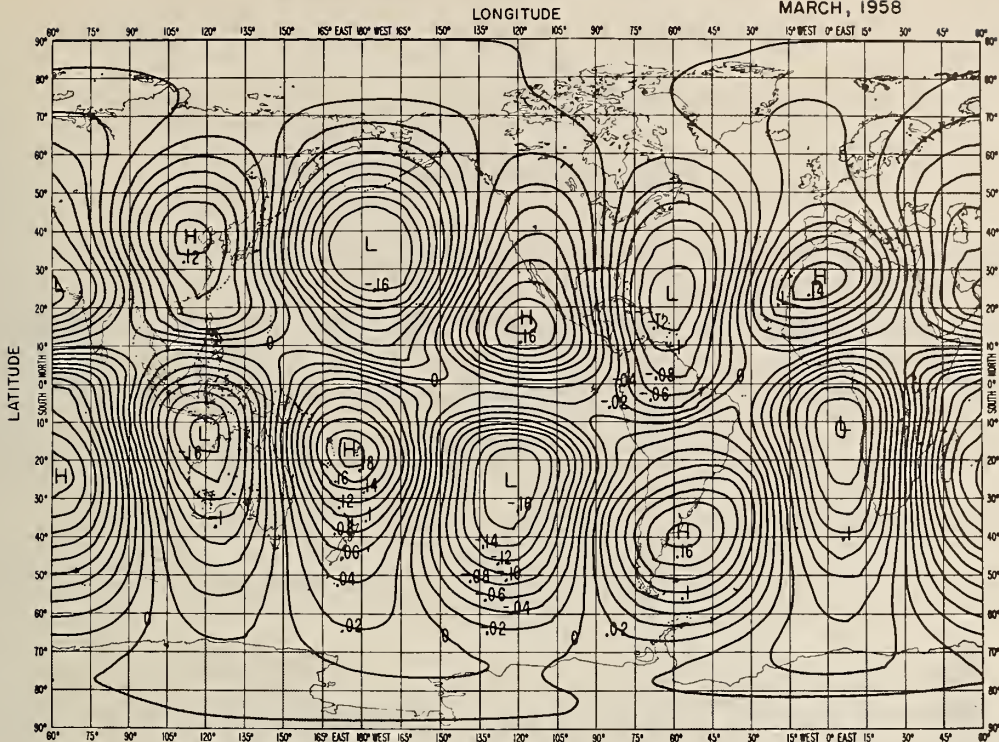
The choice of the highest degree in  $\sin x$  for each order of longitude  $\theta$  shown in Table 3 has been made from studying tests of significance of the terms  $d_k F_k(\lambda, \theta)$  in the series (11) for several typical hours in universal time from three seasonal months: March, June and December. For testing purposes, the powers of  $\sin x$  (for each order of longitude  $\theta$ ) were carried beyond those shown in Table 3, so that optimum cutoffs could be determined using the Student's t test [Jones and Gallet, 1962a]. For example, in the case of zeroth order longitude, powers of  $\sin x$  were carried out to degree 15, although degree 11 was the highest power used. Similar tests were made for the other orders of longitude. The studies showed that the optimum degree in  $\sin x$  for a given order of longitude depends upon the hour of universal time, the month and the year. The highest degrees in  $\sin x$  given by Table 3 are roughly the maxima of the optimum degrees for all hours and months tested. Thus, for a given hour and month, the series (11) may have terms which are not significant, and hence (11) may not be as efficient as possible for that particular time. This situation has been brought about intentionally, however, in order to preserve uniformity in the representations. This procedure makes it possible to analyze the individual terms in (11) from hour to hour for the diurnal analysis. It is also useful for analyzing the individual terms of numerical maps from month to month for making long term predictions.

To gain insight into the nature of the approximations of the form (11) it is instructive to examine the graphs of the orthonormal functions  $F_k(\lambda, \theta)$ , given by contour maps. As an example, figure 6 shows the functions  $F_{56}$  and  $F_{57}$ , whose leading terms ( $G_{56}$  and  $G_{57}$ ) are the third

Figure 6

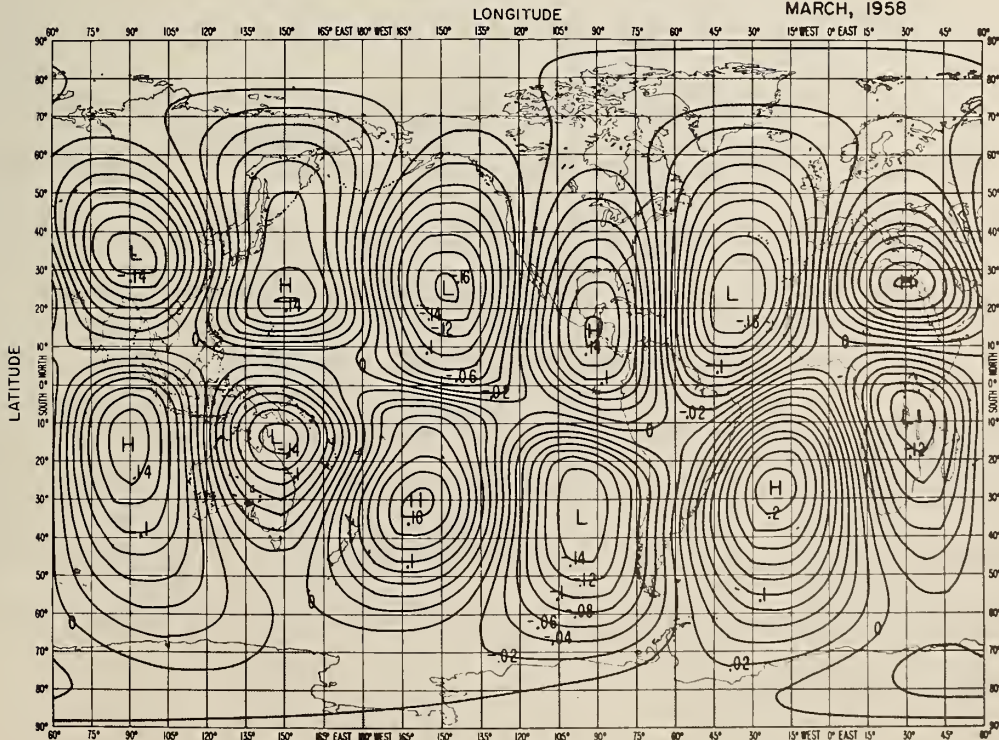
CONTOUR MAP OF ORTHONORMAL FUNCTION  $F_{56}(\lambda, \theta)$   
(HIGHEST ORDER TERM  $\sin x \cos^3 \lambda \cos 3\theta$ )

MARCH, 1958



CONTOUR MAP OF ORTHONORMAL FUNCTION  $F_{57}(\lambda, \theta)$   
(HIGHEST ORDER TERM  $\sin x \cos^3 \lambda \sin 3\theta$ )

MARCH, 1958



order longitude functions  $\sin x \cos^3 \lambda \cos 3 \theta$  and  $\sin x \cos^3 \lambda \sin 3 \theta$ , respectively. Several additional examples are contained in Appendix B. Among the properties illustrated are the close relation to magnetic dip and the position of zeros and relative maximum and minimum points.

### 3.2 Diurnal Variation

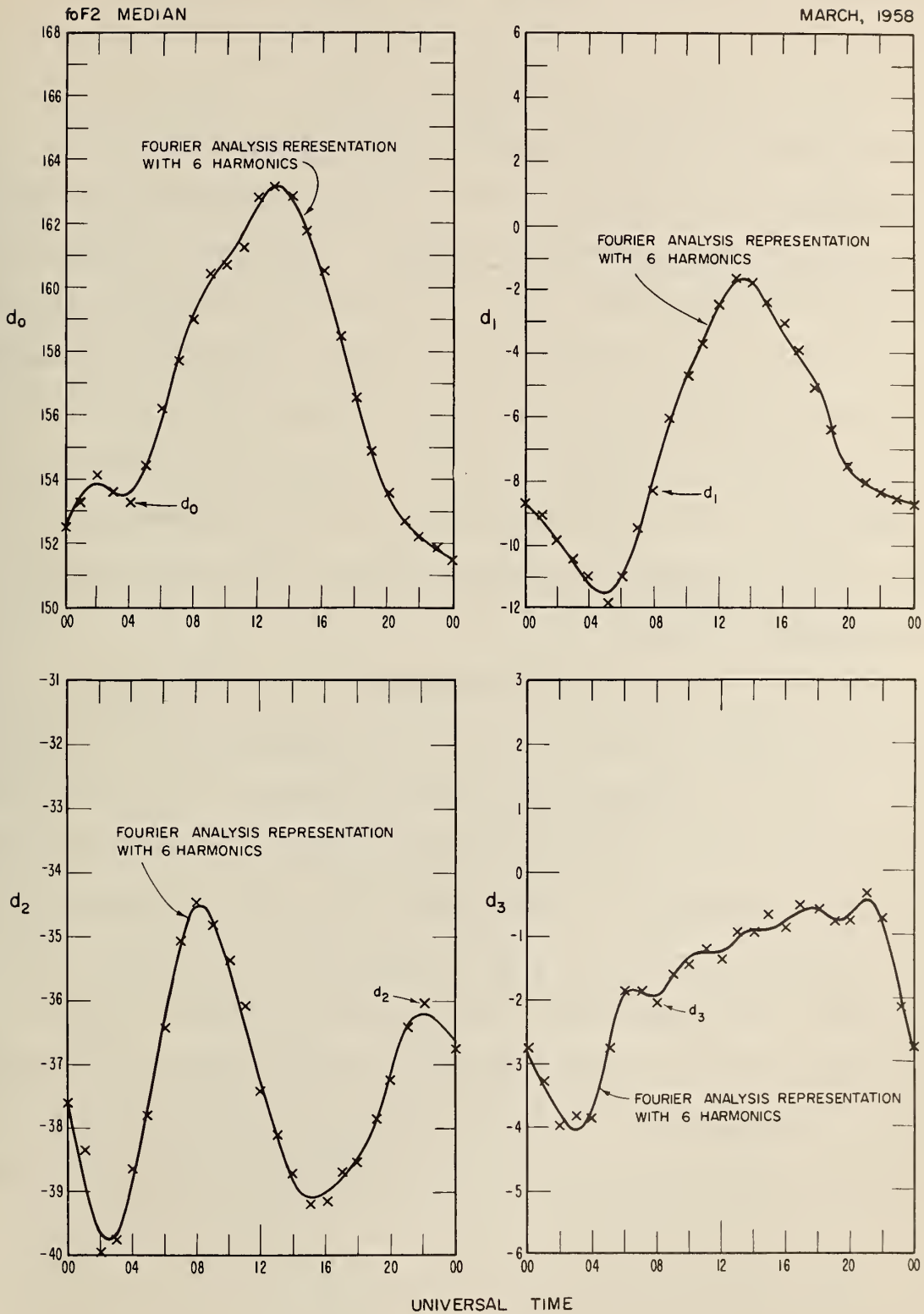
A numerical map representing both the geographic and diurnal variations of the ionospheric characteristic can now be formed by Fourier analyzing the orthonormal coefficients  $d_k$  ( $k$  fixed) obtained from the geographic analyses for the 24 hours of universal time. Typical examples of these coefficients are shown in figure 7 for  $k = 0, 1, 2, 3$  (March 1958). Also shown in these graphs are the Fourier representations of the form

$$d_k(T) = a_0^{(k)} + \sum_{j=1}^6 \left[ a_j^{(k)} \cos jT + b_j^{(k)} \sin jT \right], \quad (12)$$

where  $T$  denotes the universal time hour angle ( $-180^\circ \leq T \leq 180^\circ$ ,  $T = 0^\circ$  at 12 noon U.T.). Further examples of this type are given in Appendix C. Only six harmonics are used for this representation since it was known from previous work [Jones, 1962; Jones and Gallet, 1962a and 1965b] that at most eight harmonics are needed for the diurnal variation and since experiments made in this application showed that harmonics seven and eight gave no significant contribution. The  $d_k$  have a remarkable degree of smoothness and continuity over a 24 hour period, in spite of the fact that the coefficients for each hour are obtained independently. Further studies have shown that the diurnal variations of the coefficients  $D_k$  in (8) are not as smooth as for the  $d_k$ .

Figure 7

DIURNAL VARIATION OF ORTHONORMAL COEFFICIENTS  $d_k$



The numerical map of the ionospheric characteristic can now be written in the form

$$\Omega(\lambda, \theta, T) = \sum_{k=0}^{75} d_k(T) F_k(\lambda, \theta), \quad (13)$$

where the functions  $d_k(T)$  have the form (12). For use in practical applications,  $\Omega(\lambda, \theta, T)$  is expressed in terms of the simpler functions  $G_k(\lambda, \theta)$  in Section 5. Examples of graphical contour maps of foF2 median for March 1958 computed from  $\Omega(\lambda, \theta, T)$  are shown in figures 8 and 9, for the hours U.T. = 00, 06, 12, 18. As can be seen, the problem of ambiguous polar values is no longer present. Also, by comparison with graphical contour maps from former analyses such as Jones and Gallet, [1965a], it is evident that the new maps represent more detail in the equatorial region. Further comparison of different analyses will be dealt with in the following section.

To conclude this section we shall mention a few of the strong connections between the various steps of the mapping procedures described above. The use of a universal time analysis eliminated polar ambiguities but, at the same time, created another difficulty: increased complication of geographic variations. To take account of that fact, it was necessary to use a larger number of geographic functions (including higher orders of longitude). This placed even more importance on C-data to control instability; hence more screen points had to be used. Since the geographic analysis is far more difficult than the diurnal, it was decided to analyze the geographic variations of the A-, B- and C- data for fixed hours of UT,

rather than the Fourier coefficients from the diurnal analysis, adjusted to UT. The latter approach, though perhaps more desirable from the viewpoint of reducing the overall number of terms needed to define a numerical map, would be more complicated since the Fourier coefficients have extremely different types of variations from one coefficient to the next. Hence a careful study would be required for each (of approximately 17) Fourier coefficient. Perhaps this is an area worthy of further investigation.

Figure 8

MAPS OF foF2 MEDIAN (Mc/s) COMPUTED FROM UT ANALYSIS

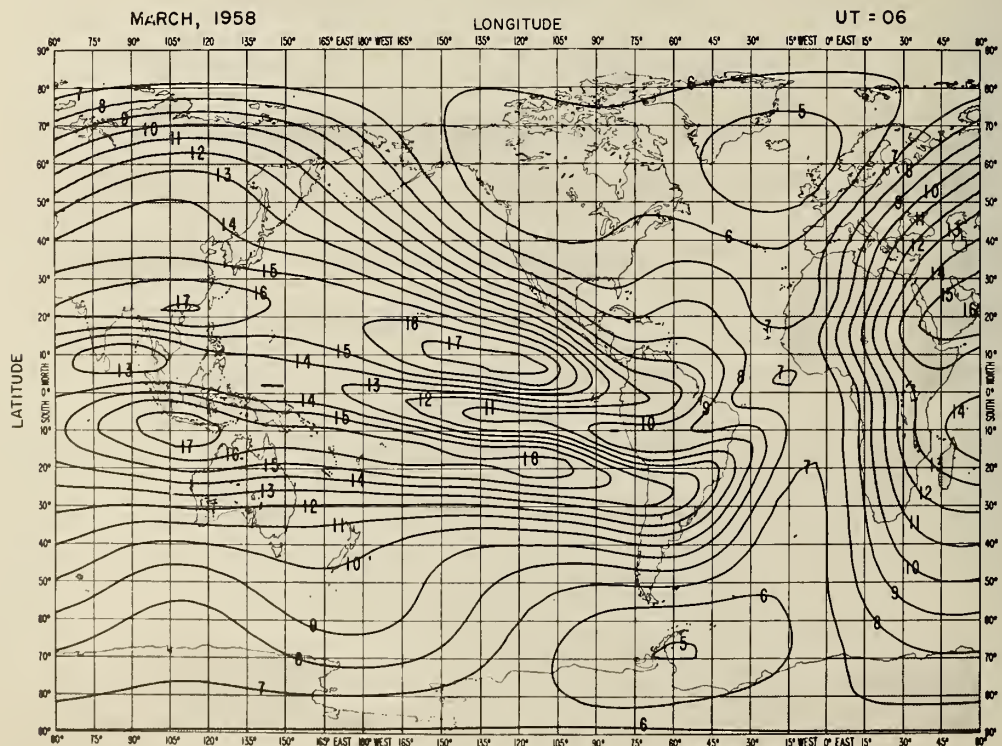
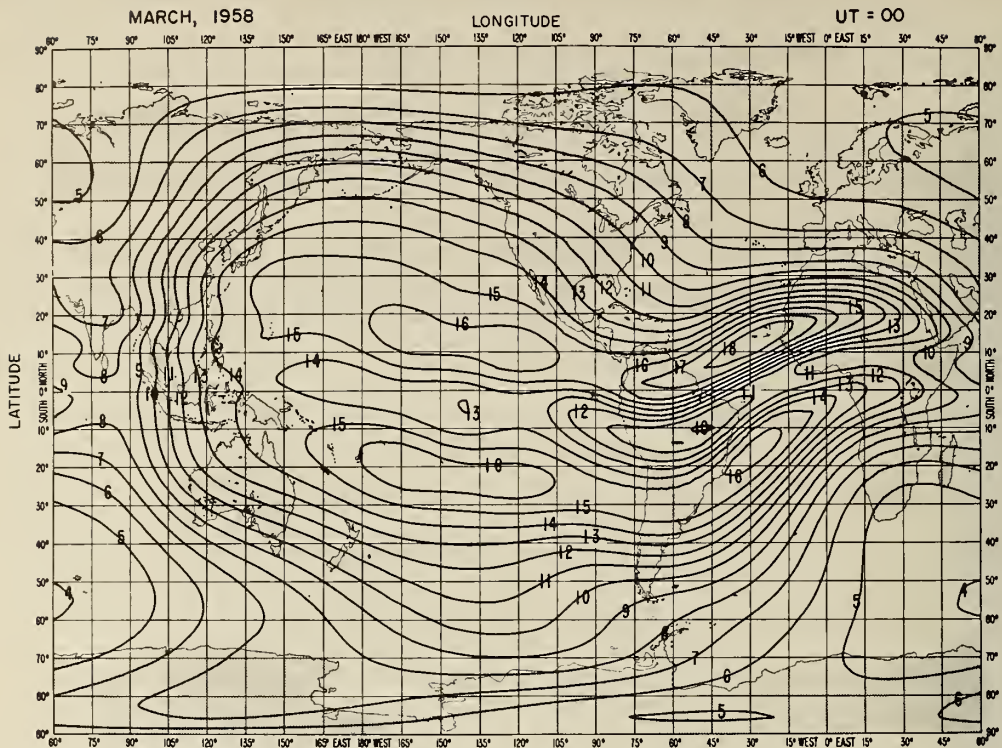
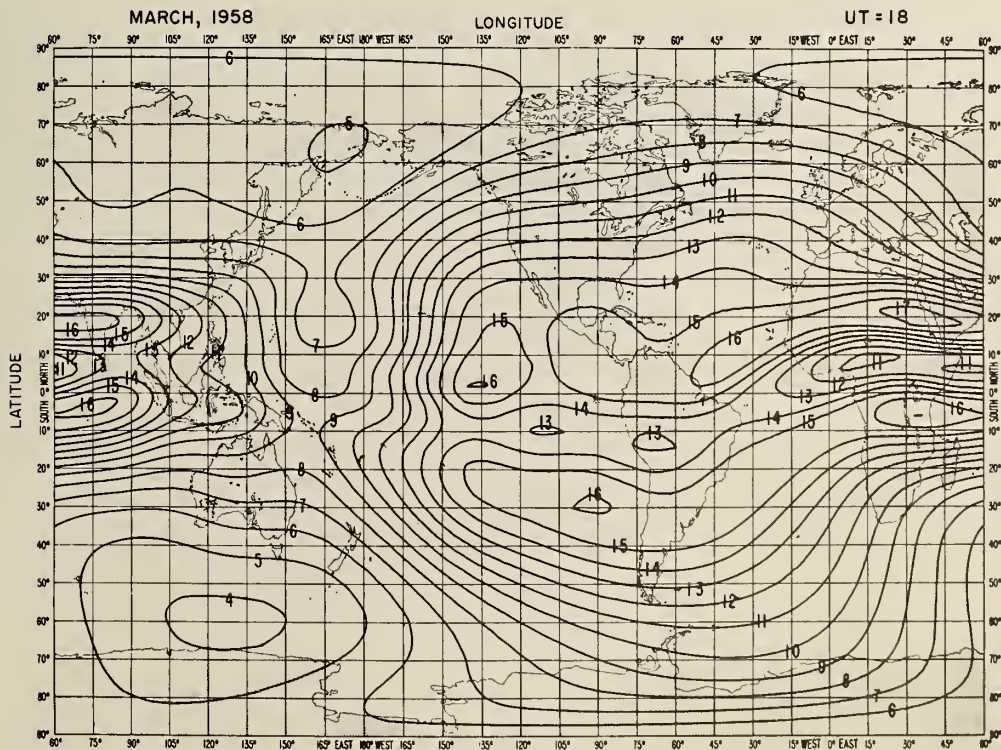
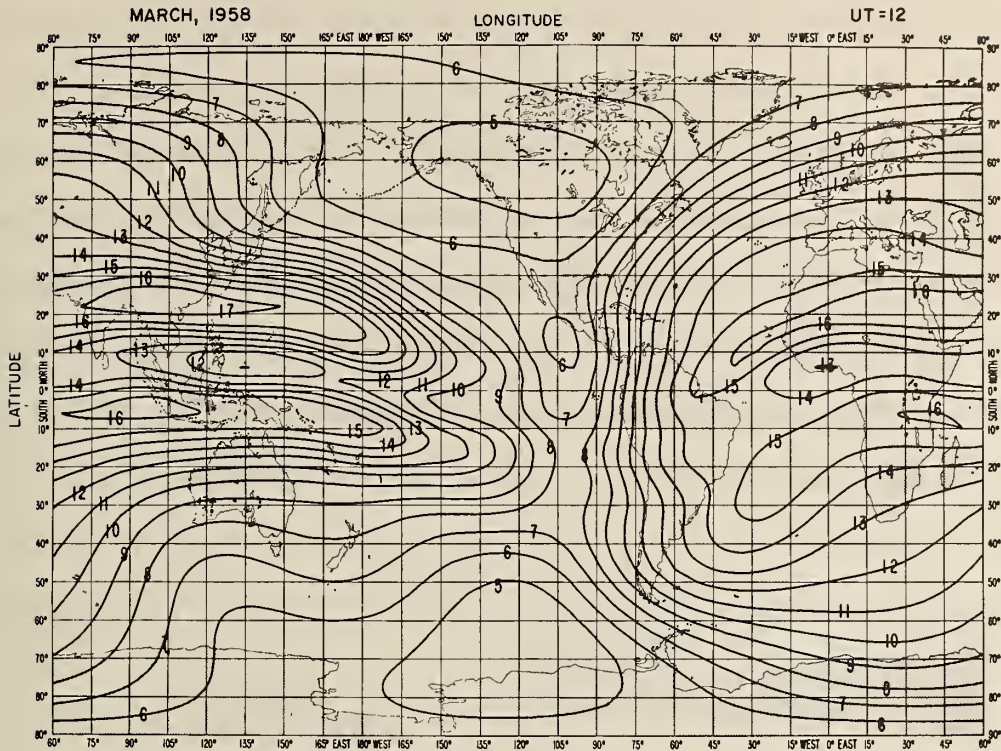




Figure 9

MAPS OF foF2 MEDIAN (Mc/s) COMPUTED FROM UT ANALYSIS



#### 4. Comparison of Different Analyses

Results from two different types of analysis are compared in this section. Type I is described by Jones and Gallet [1965a]. Local mean time is used for both the screen and second analyses. The geographic variable for the screen analysis is latitude  $\lambda$ . For the second analysis the geographic functions are similar to those given in Table 3, with a maximum of two orders of longitude  $\theta$  and with the factor  $\sin x$  replaced by  $\sin \lambda$ . Type II is described in the present paper (Sections 2 and 3). Local mean time is used in the screen analysis and universal time in the second analysis. The geographic variable for the screen analysis is the modified magnetic dip  $x$ , see (4), and the geographic functions for the second analysis are given in Table 3. Table 4 summarizes the main characteristics of the two types of analysis.

Table 4

Main Characteristics of Two Types of Analysis

	Screen Analysis		Second Analysis	
	Time	Geographic Variable	Time	Geographic Variables
Type I	LMT	$\lambda = \text{latitude}$	LMT	$\lambda$ and $\theta$
Type II	LMT	$x = \text{modified magnetic dip}$	UT	$\lambda$ , $\theta$ , and $x$

( $\theta = \text{geographic longitude}$ )

For comparison, the two types of analysis have been applied to the same sets of A- and B-data for the characteristic foF2 median for four seasonal months of the high and low years of solar activity, 1958 and 1954. For each type analysis and each month, the root mean square (RMS) residual has been computed using residuals between A- and B-data and numerical map values for all stations and all 24 hours. The values of the RMS residual, which estimate the standard deviation of residuals, are given in Table 5. As can be seen, there is a major improvement in the representation of the given data. The reduction in the RMS residual for Type II ranges from 30 to 47 per cent. The values of the RMS residual

Table 5

RMS of Residuals of (Mc/s) of A- and B-data for Two Types of Analysis  
(All hours and all stations of foF2 medians)

Analysis	1954				1958			
	Mar.	June	Sept.	Dec.	Mar.	June	Sept.	Dec.
Type I	.72	.58	.69	.66	1.10	1.06	1.09	.98
Type II	.38	.36	.41	.42	.63	.59	.60	.67

in Table 5 for Type I are considerably higher than those reported in the previous paper [Jones and Gallet, 1965a]. This is due to the use of about 30 additional B-data stations. The number of stations with A- or B-data and the number of screen points for each month used in the present study are summarized in Table 6.

Table 6

Number of Stations ( $N_A$  and  $N_B$ ) with A- and B-data and  
 Number of Screen Points ( $N_C$ ) used in Analysis of foF2 Median

	1954				1958			
	Mar.	June	Sept.	Dec.	Mar.	June	Sept.	Dec.
$N_A$	87	87	85	83	121	124	124	122
$N_B$	71	71	75	76	42	38	38	41
$N_C$	68	68	68	68	67	67	67	67

By comparing the results of Table 5 for Type II analysis with estimates of the noise inherent in foF2 medians [Jones, 1962; Jones and Gallet, 1962a, 1965b], it appears that the residuals in the numerical map representations are now close to the noise level. Although some additional geographic functions may be required to fully represent the physical variations in existing data, it is felt that a further extension of the series (11) could not be justified with the presently available distribution of ionospheric sounding stations. To summarize, the Type II analysis is superior to Type I on three major counts: (1) closer fitting of the given A- and B-data, (2) more accurate representation of the ionospheric characteristic in critical areas, a result of having improved C-data and (3) elimination of ambiguous values at the geographic poles.

## 5. Application of Numerical Maps

### 5.1 Simplified Expression of $\Omega(\lambda, \theta, T)$

As an aid to evaluating numerical maps, this section reduces the function  $\Omega(\lambda, \theta, T)$  to a form expressed in terms of the geographic functions  $G_k(\lambda, \theta)$ . Also included in this section are the formulas necessary for calculating magnetic dip I. To obtain increased generality, so that the discussion will apply to numerical maps of any ionospheric characteristic, we introduce a set of geographic functions  $G_k(\lambda, \theta)$  in Table 7, of which Table 3 is a special case. Here the integer  $q_i$ ,  $i = 0, 1, \dots, m$ , denotes the highest power of  $\sin x$  for the  $i^{\text{th}}$  order harmonic in longitude, and  $m$  denotes the highest order of longitude. The parameters  $q_i$  and  $m$  may change from one ionospheric characteristic to another. However, the general form of the numerical maps will remain fixed. For indexing purposes it is sometimes convenient to introduce the sequence  $k_i$ ,  $i = 0, 1, \dots, m$ , such that  $G_{k_i}(\lambda, \theta)$  is the last of the geographic functions of  $i^{\text{th}}$  order longitude:

$$G_{k_0}(\lambda, \theta) = \sin^{q_0} x$$

$$G_{k_i}(\lambda, \theta) = \sin^{q_i} x \cos^i \lambda \sin i\theta, \quad i = 1, 2, \dots, m. \quad (14)$$

Thus

$$q_0 = k_0; \quad q_i = \frac{k_i - k_{i-1}}{2}, \quad i = 1, 2, \dots, m, \quad (15)$$

and the total number of geographic functions in Table 7 is given by

$$K + 1 = k_m + 1 .$$

For convenience in computing procedures we note that

$$\sin x = \frac{I}{\sqrt{I^2 + \cos \lambda}} \quad (16)$$

which follows immediately from (4).

Table 7

Geographic Coordinate Functions

$G_k(\lambda, \theta)$  ( $x$  is a function of  $\lambda, \theta$ ),  $m$  is the maximum order in longitude.

$k$	Main Lat. Variation	$k$	First Order Longitude $G_k$	$k$	Second Order Longitude $G_k$	...	$k$	$m$ th Order Longitude
0	1	$k_0+1$	$\cos \lambda \cos \theta$	$k_1+1$	$\cos^2 \lambda \cos 2 \theta$	...	$k_{m-1}+1$	$\cos^m \lambda \cos m \theta$
1	$\sin x$	$k_0+2$	$\cos \lambda \sin \theta$	$k_1+2$	$\cos^2 \lambda \sin 2 \theta$	...	$k_{m-1}+2$	$\cos^m \lambda \sin m \theta$
2	$\sin^2 x$	$k_0+3$	$\sin x \cos \lambda \cos \theta$	$k_1+3$	$\sin x \cos^2 \lambda \cos 2 \theta$	...	$k_{m-1}+3$	$\sin x \cos^m \lambda \cos m \theta$
.	.	$k_0+4$	$\sin x \cos \lambda \sin \theta$	$k_1+4$	$\sin x \cos^2 \lambda \sin 2 \theta$	...	$k_{m-1}+4$	$\sin x \cos^m \lambda \sin m \theta$
.	.	.	.	.	.	.	.	.
.	.	.	.	.	.	.	.	.
$k_0$	$\sin^{q_0} x$	$k_1-1$	$\sin^{q_1} x \cos \lambda \cos \theta$	$k_2-1$	$\sin^{q_2} x \cos^2 \lambda \cos 2 \theta$	...	$k_m-1$	$\sin^{q_m} x \cos^m \lambda \cos m \theta$
		$k_1$	$\sin^{q_1} x \cos \lambda \sin \theta$	$k_2$	$\sin^{q_2} x \cos^2 \lambda \sin 2 \theta$	...	$k_m$	$\sin^{q_m} x \cos^m \lambda \sin m \theta$

A numerical map, as described in Section 3, may then be written in the general form

$$\Omega(\lambda, \theta, T) = \sum_{k=0}^K d_k(T) F_k(\lambda, \theta), \quad (17)$$

where

$$d_k(T) = a_0^{(k)} + \sum_{j=1}^H \left[ a_j^{(k)} \cos jT + b_j^{(k)} \sin jT \right], \quad (18)$$

where  $H$  is the number of harmonics retained for representing the diurnal variation of the orthonormal coefficients  $d_k$ . Here  $F_k(\lambda, \theta)$  denotes the

orthonormal functions, discussed in Section 3, which are linear combinations of the  $G_k(\lambda, \theta)$  in the form (9). In most applications it is convenient to express  $\Omega(\lambda, \theta, T)$  explicitly in terms of the  $G_k(\lambda, \theta)$ . If (18) is substituted in (17) and the order of summation interchanged, we obtain

$$\Omega(\lambda, \theta, T) = a_0(\lambda, \theta) + \sum_{j=1}^H \left[ a_j(\lambda, \theta) \cos jT + b_j(\lambda, \theta) \sin jT \right], \quad (19)$$

where

$$a_j(\lambda, \theta) = \sum_{k=0}^K a_j^{(k)} F_k(\lambda, \theta), \quad (20)$$

$$b_j(\lambda, \theta) = \sum_{k=0}^K b_j^{(k)} F_k(\lambda, \theta).$$

Now the procedure described by Jones and Gallet [1962a, Section 2.2] may be used to obtain the coefficients  $U_{s,k}$  for defining the  $a_j(\lambda, \theta)$  and  $b_j(\lambda, \theta)$  in the form

$$a_j(\lambda, \theta) = \sum_{k=0}^K U_{2j,k} G_k(\lambda, \theta), \quad j = 0, 1, \dots, H, \quad (21)$$

$$b_j(\lambda, \theta) = \sum_{k=0}^K U_{2j-1,k} G_k(\lambda, \theta), \quad j = 1, 2, \dots, H. \quad (22)$$

Thus a numerical map  $\Omega(\lambda, \theta, T)$  is defined by specifying the integers  $q_0, q_1, \dots, q_m, H$ , and the coefficients  $U_{s,k}$  in (21) and (22). As an example, for the numerical map of foF2 median for March 1958 obtained from the Type II analysis in Section 3, we have  $H = 6$ ,  $q_0 = q_1 = 11$ ,  $q_2 = 8$ ,  $q_3 = 4$ ,  $q_4 = 1$ ,  $q_5 = q_6 = q_7 = q_8 = 0$ . The coefficients  $U_{s,k}$  are given in Table 8.





Table 8 (contd.)

TIME VARIATION

Harmonic	4				5				6				
	7	8	9	10	11	12	13	14	15	16	17	18	
I	0	-1.0779941-002	-2.3504205-002	5.0223169-002	-1.5342471-002	-8.1190315-003	1.457496A-002						
	1	-2.5163831-001	3.4951377-004	1.2539203-001	-2.927617A-002	4.518A227-002	-6.1192357-002						
	2	1.7617797-001	-3.0042594-001	-1.9069479-000	4.9804011-001	1.2314179-001	-1.3595216-001						
	3	5.0009802-000	2.1337716+000	-1.0721361-000	2.3865944-001	-2.4400391-001	1.4734534+000						
	4	1.4481989+000	4.9181347+000	1.1474365+001	-4.2171784+000	-1.1915910-001	-3.0943748+000						
	5	-7.7024190+001	-2.0017166+001	4.6253478+000	-5.3854807-001	-1.3510171+000	-8.9562380+000						
	6	-8.0176203+000	-1.6743417+001	-2.726A753+001	1.2287A97+000	-8.8804934-001	2.779A981+000						
	7	6.0793524+001	5.75A2418+001	-9.415217A+000	-5.3474611-001	-4.0530766+000	2.1151A46+001						
	8	1.1556410+001	2.1265702+001	2.8373627+001	-1.4409546+001	1.9977780+000	-3.2A50628+000						
	9	-6.0950502+001	-6.5751104+001	A.3870A50+000	2.2717A81+000	5.4099207+000	-7.1300741+001						
	10	-5.0951467+000	-9.2074405+000	-1.0743671+001	5.8734569+000	-1.1181323+000	1.44A9519+000						
11	2.2507631+001	2.6079102+001	-2.6005567+000	-1.4104205+000	-2.4713309+000	7.6567A81+000							
II	12	-1.1458245-002	-6.9024523-002	2.7850092-002	2.0A72289+002	-1.7520945-002	-2.9.001496+004						
	13	-2.2159441-002	5.8005822-002	6.1969955-002	2.02762A2+002	-A.9692914+003	-2.2068646-002						
	14	1.6053418-002	2.9000543-001	9.8697105-002	-8.6A13793+003	-1.4062745-001	-1.488A9836-002						
	15	3.5166447-001	-3.3348892-002	1.657567A-001	3.2135004-003	-8.6025890-002	-1.3032900-001						
	16	1.7864647+000	1.345A96A+000	-1.0224534+000	-2.73511A1-001	5.8111404-001	-1.063A485+001						
	17	2.9747778+000	-2.8411192+001	-3.3289101+000	-2.5100550+001	-1.9230900+001	3.4427307+001						
	18	1.0961566+000	-4.5654592+000	-5.6R44113-001	1.762A557-001	3.1364766+000	3.4254672-001						
	19	-7.1431932+000	1.0872520+000	-3.4736647+000	-1.53507A1+000	1.2499952+000	1.520295A+000						
	20	-1.0806991+001	-9.5641344+000	6.27327A1+000	3.1002005+000	-4.2654736+000	2.0729756+000						
	21	-2.2218944+001	5.924A843+000	2.863A919+001	9.5760513+001	2.305A672+000	-1.3706267+000						
	22	-1.4468506+001	2.0055106+001	-2.346A676+000	-7.0226575+000	-1.8265760+001	-3.1887424+000						
23	3.7685009+001	2.4442954+000	1.6178771+001	1.4441918+001	-9.1065089+000	-7.1897774+000							
37	-2.6693593+001	2.8433531+001	-1.6216755+001	-1.1933240+001	1.2144399+001	-8.74408331+000							
25	5.6381110+001	-8.2394477+000	-8.3706335+000	-2.2815355+000	-8.72A1207+000	2.4226267+000							
26	4.8174284+000	-3.9401106+001	1.6153972+001	2.755A36A+001	4.2961625+001	1.250717A+001							
27	-7.1530434+001	-3.0154509+001	-2.9144632+001	-4.4405647+001	2.38A357+001	1.9211623+001							
28	-3.0414312+001	-3.6674649+001	1.9500457+001	1.7255337+001	-1.4543346+001	1.28A3081+001							
29	-5.5445028+001	1.8454919+000	1.0006330+002	5.3911944+000	1.324A117+001	-2.0610975+000							
30	-6.2982974+001	3.64A1757+001	-2.6153372+001	-3.6670009+001	-4.4244550+001	-1.9667104+001							
31	4.8136537+001	5.8187436+001	2.1969150+001	5.602214A+000	-2.5961557+001	-2.6A93497+001							
32	1.3081566+001	1.6511594+001	-8.8923541+000	-8.5467240+001	6.1793288+000	-6.1720834+000							
33	1.7281032+001	1.0444953+000	-4.2244317+001	-2.8479104+000	-6.9110015+000	7.170A903+001							
34	2.9115728+001	-1.3607991+001	1.3121563+001	1.60419A7+000	1.6694152+001	1.044A565+001							
35	-5.7052288+000	-3.0149413+001	-5.5163518+000	-2.5754167+001	9.33A1503+000	1.4280357+001							
III	36	1.6620753+002	9.350A189+003	-1.7836453+002	2.8060952+002	3.801A123+002	-5.2266452+002						
	37	-6.6120614+003	2.3405097+002	1.9616151+002	2.6330532+002	-2.1868669+003	-8.6431694+003						
	38	-2.0609766+001	-1.4721618+001	-6.4175082+002	7.3481436+002	-1.5299020+002	-1.0959351+001						
	39	-6.3109410+001	1.8A21194+001	1.2855597+001	-1.3786A97+001	-7.695A922+002	5.3034327+002						
	40	1.1615553+000	2.2402236+001	3.8100A05+001	-1.1631414+001	-9.2070603+001	1.2175463+000						
	41	1.5621639+000	-7.0442700+001	-1.4162618+000	-5.1350730+001	2.030A040+001	-4.3672179+002						
	42	1.8278950+000	2.4269778+000	3.9492903+001	-1.2481072+000	1.2200315+000	1.397360A+000						
	43	3.5461865+000	-4.4384758+001	-1.4466332+000	1.23964A9+000	-3.83359A8+002	-2.17302A2-001						
	44	3.6879955+000	-3.8096056+000	-7.853A335+001	-1.6987433+001	5.0403690+000	-7.1650A4A+000						
	45	-1.0388171+001	1.1114892+000	5.8923541+000	1.6490190+000	-6.0164055+001	-1.0906841+000						
	46	-4.6440333+000	-8.7567035+000	-7.8604977+001	3.8512794+000	-2.6309910+000	-3.78A6949+000						
47	-6.9636327+000	4.2168317+001	2.9414559+000	-3.4211997+000	6.4303710+001	2.67521A8+001							
48	2.7142404+001	1.2443316+001	-6.6582155+001	1.50A2624+000	-9.36A7013+000	1.48A8593+001							
49	2.0402A81+001	8.0A8A18A+001	-2.1219091+001	-1.2770435+000	5.119A332+001	2.8670662+000							
50	3.6660034+000	A.2400076+000	7.3676752+001	-2.9591239+000	1.6183555+000	2.6943651+000							
51	3.8A9A490+000	-7.2737664+001	-1.7140A31+000	2.96A0919+000	-4.8090416+001	-3.042930A+002							
52	-1.4403383+001	-1.1164680+001	1.4823537+000	-1.4756282+000	5.4644740+000	-9.8710287+000							
53	-1.2431180+001	-1.07A78A2+000	1.4460124+001	-2.6742897+001	-2.2700363+001	-2.5799779+000							
IV	54	7.2832114+002	-2.03A9792+002	7.8724A66+003	2.8061195+002	2.0920550+002	-4.4055183+002						
	55	1.523921A+002	-4.9400445+003	-1.6769639+002	-9.2680243+003	A.761A165+003	5.149AA6+003						
	56	4.4816603+002	2.7177493+001	1.7797035+002	-5.4446760+002	9.2971316+002	-2.2A25522-002						
	57	-2.2603405+001	1.0007460+001	3.0096395+002	8.5965596+003	-6.0280A55+004	4.7255015+002						
	58	-4.4921730+001	-1.2482743+001	7.9967042+002	-5.3480174+002	-1.1994559+001	4.588003A+001						
	59	3.5216262+002	7.5449431+002	-6.9711364+002	2.0900151+001	3.1421447+002	-2.3092020+001						
	60	-4.8451815+003	-4.7801694+001	-5.0020294+002	2.6527609+001	-2.7050890+001	7.8751A6A+002						
V	61	4.9907794+001	-6.930A474+002	-8.1722447+002	8.243012A+002	-3.65667A1+002	-6.3659351+002						
	62	5.7205307+001	3.7554418+001	-6.1636232+002	-6.8014707+002	3.2620970+001	-6.6174797+001						
	63	-2.5666734+001	-1.4794603+001	3.1409364+001	-4.3419296+001	-6.2514650+002	1.823A921+001						
	64	2.8461028+001	-3.944147+001	2.015757A+002	2.8399051+003	-1.1443148+003	-A.0223370+003						
VI	65	3.5799998+001	2.6751436+001	-3.0073297+002	-6.7519533+003	9.7485471+003	-6.78A4402-003						
	66	1.3543195+001	-1.6540876+001	-1.1847066+001	7.7174367+002	1.9120149+003	3.215A505+002						
	67	1.1650050+001	1.1836679+001	-2.555264A+002	-2.871A601+002	2.400397A+003	1.5924910+002						
VII	68	-4.2422449+002	-3.3732559+002	2.9134A23+001	1.6649069+001	-9.4621915+004	-2.1111944+003						
	69	-3.6958211+003	6.7456529+003	-1.8252193+001	2.6141323+001	-9.313A324+003	2.0207452+002						
VIII	70	-4.1576270+002	-A.424A8A0+003	1.8739525+002	1.021493A+002	-1.0852350+001	1.5656743+001						
	71	6.5759298+003	-3.264A854+003	-9.1919265+004	A.9791466+003	-1.8366006+001	-7.307436A+002						
IX	72	-1.6174064+002	-6.074042A+003	5.6854444+003	-2.8063166+002	-2.4771362+003	6.4377171+003						
	73	6.2887233+003	3.4130607+003	5.0504742+004	2.8553642+002	-7.4174002+003	1.1032996+003						
X	74	4.61A2573+002	1.5727465+002	3.1559A9A+002	-3.2167669+002	2.1A57576+002	1.03455A0+002						
	75	4.7004174+002	-2.2495015+002	-1.4974033+002	1.6224500+002	-1.62A8146+002	-2.04378A9+003						

I - Main latitudinal variation. Mixed latitudinal and longitudinal variation; II - First order in longitude;

III - Second order in longitude. .... IX - Eighth order in longitude.

Notation: For each entry the number given by the first eight digits and sign is multiplied by the power of ten defined by the last three digits and sign.

COEFFICIENTS  $U_{SK}$  DEFINING THE FUNCTION  $\Omega(\lambda, \theta, T)$  FOR  
MONTHLY MEDIAN  $f_0F2$  (Mc/s)  
MARCH 1958

## 5.2 Calculation of Magnetic Dip I

To evaluate the geographic functions  $G_k(\lambda, \theta)$  in Table 7 for given values of  $\lambda$  and  $\theta$ , it is necessary to calculate magnetic dip I, from which we may obtain the modified magnetic dip  $x$  defined by (4). We include here for completeness the necessary formulas. Introduce the notation for colatitude,  $\varphi = 90^\circ - \lambda$ , and

$$R = \frac{a}{a + h} \quad (23)$$

where  $a = 6371.2$  km is the radius of the earth and  $h$  is height above sea level in kilometers. For F2 layer characteristics the value  $h = 300$  km has been used. Thus we compute

$$I = \text{Tan}^{-1} \left[ \frac{-Z}{\sqrt{X^2 + Y^2}} \right], \quad (24)$$

where  $X$ ,  $Y$ , and  $Z$  are, respectively, the north, east, and vertical (down) components of the magnetic field vector  $\underline{F}$ . Following the discussion of spherical harmonic analysis of the magnetic field by Chapman and Bartels [1962], we have

$$X = \sum_{n=1}^6 \sum_{m=0}^n x_n^m \left[ g_n^m \cos m \theta + h_n^m \sin m \theta \right] R^{n+2}, \quad (25)$$

$$Y = \sum_{n=1}^6 \sum_{m=0}^n y_n^m \left[ g_n^m \sin m \theta - h_n^m \cos m \theta \right] R^{n+2}, \quad (26)$$

$$Z = \sum_{n=1}^6 \sum_{m=0}^n z_n^m \left[ g_n^m \cos m \theta + h_n^m \sin m \theta \right] R^{n+2}, \quad (27)$$

where

$$x_n^m = \frac{d}{d\varphi} P_{n,m}(\cos \varphi), \quad (28)$$

$$y_n^m = \frac{m P_{n,m}(\cos \varphi)}{\sin \varphi} \quad (29)$$

$$z_n^m = -(n+1) P_{n,m}(\cos \varphi). \quad (30)$$

Here the function  $P_{n,m}(\cos \varphi)$  is a multiple of the associated Legendre function and is defined by

$$P_{n,m}(\cos \varphi) = \sin^m \varphi \left[ \cos^{n-m} \varphi - \frac{(n-m)(n-m-1)}{2(2n-1)} \cos^{n-m-2} \varphi \right. \\ \left. + \frac{(n-m)(n-m-1)(n-m-3)}{(2)(4)(2n-1)(2n-3)} \cos^{n-m-4} \varphi - \dots \right]. \quad (31)$$

The coefficients  $g_n^m$  and  $h_n^m$  used in the present work are those given by Jensen and Cain [1962] from the analysis of the magnetic field for Epoch 1960.

### 5.3 Fortran Program

Appendix D contains a FORTRAN program which may be used to evaluate a numerical map  $\Omega(\lambda, \theta, T)$  of an ionospheric characteristic such as foF2 median, or to calculate the F2 layer maximum usable frequencies for zero and 4000 kilometers [MUF(ZERO)F2 and MUF(4000)F2]. We introduce the following notation:

$$(1) \quad \lambda = \text{geographic latitude,} \quad -90^\circ \leq \lambda \leq 90^\circ,$$

$$(2) \quad \theta = \text{east geographic longitude} \quad 0^\circ \leq \theta \leq 360^\circ,$$

$$(3) \quad h_\lambda = \text{hours in local mean time (LMT)} \quad 0 \leq h_\lambda \leq 24,$$

$$(4) \quad h_u = \text{hours in universal time (UT)} \quad 0 \leq h_u \leq 24,$$

$$(5) \quad T = \text{universal time hour angle,} \quad -180^\circ \leq T \leq 180^\circ,$$

$$T = 15 h_u - 180 = 15 h_\lambda + 180 - \theta$$

$$(6) \quad B = \text{intensity of magnetic field,}$$

$$(7) \quad f_H = \text{gyrofrequency}$$

The following three examples are given for the use of the FORTRAN program.

Example I. Compute foF2 median (or any other ionospheric characteristic) for a fixed instant  $h_u$  of universal time at several geographic points  $(\lambda, \theta)$ .<sup>1</sup>

$$(a) \quad \text{Compute } T_o = 15 h_u - 180.$$

$$(b) \quad \text{Evaluate } \cos jT_o \text{ and } \sin jT_o, \quad j = 1, 2, \dots, H.$$

$$(c) \quad \text{Compute } D_k(T_o) = U_{o,k} + \sum_{j=1}^H \left[ U_{2j,k} \cos jT_o + U_{2j-1,k} \sin jT_o \right],$$

$k = 0, 1, \dots, K$ , (where the  $U_{s,k}$  are given either by a table such as Table 8 or on cards or tape.

<sup>1</sup>The subscript o indicates a fixed value of the parameter.

- (d) Compute magnetic field components X, Y, Z for the given point  $(\lambda, \theta)$  and height = 300 km

(Use Subroutine MAGFIN in program SAMPLE.),

$$I = \tan^{-1} \left( -Z / \sqrt{X^2 + Y^2} \right),$$

$$\sin x = I / \sqrt{I^2 + \cos \lambda}.$$

- (e) Evaluate  $G_k(\lambda, \theta)$   $k = 0, 1, \dots, K$ .

$$(f) \text{ Compute } \Omega_F(\lambda, \theta, T_0) = \sum_{k=0}^K D_k(T_0) G_k(\lambda, \theta).$$

Repeat steps d, e, f for each given point  $(\lambda, \theta)$ ; steps a, b, c need not be repeated if  $T_0$  remains fixed. Here  $\Omega_F$  is used to denote the numerical map value of foF2 median (Mc/s). Similarly  $\Omega_m$  will be used later to denote the map value of the characteristic M(3000)F2 median.

Example II. Compute MUF(ZERO)F2 at a fixed geographic point

$(\lambda_0, \theta_0)$  for several instants  $h_u$  of universal time.

- (a) Compute X, Y, Z, I and  $\sin x$  for the fixed point  $(\lambda_0, \theta_0)$  and height = 300 km as in Example I, step (d).

- (b) Compute  $B = \sqrt{X^2 + Y^2 + Z^2}$  and  $f_H = 2.8 B$ .

- (c) Evaluate  $G_k(\lambda_0, \theta_0)$ ,  $k = 0, 1, \dots, K$  (for map of foF2 median).

$$(d) \text{ Compute } a_j(\lambda_0, \theta_0) = \sum_{k=0}^K U_{2j,k} G_k(\lambda_0, \theta_0), \quad j = 0, 1, \dots, H,$$

$$b_j(\lambda_0, \theta_0) = \sum_{k=0}^K U_{2j-1,k} G_k(\lambda_0, \theta_0), \quad j = 1, 2, \dots, H,$$

using coefficients  $U_{s,k}$  for map of foF2 median.

(e) Compute  $T = 15 h_u - 180$ .

(f) Evaluate  $\cos jT$  and  $\sin jT$ ,  $j = 1, 2, \dots, H$ .

(g) Compute  $\Omega_f(\lambda_o, \theta_o, T) = a_o(\lambda_o, \theta_o) + \sum_{j=1}^H \left[ a_j(\lambda_o, \theta_o) \cos jT + b_j(\lambda_o, \theta_o) \sin jT \right]$ .

(h)  $MUF(ZERO)F2 = \Omega_f + f_H/2$ .

For each  $T$  repeat steps e - h; steps a - d need not be repeated if  $(\lambda_o, \theta_o)$  remains fixed.

Example III. Compute  $MUF(4000)F2$  for a given point  $(\lambda_o, \theta_o)$  for several instants  $h_\ell$  of local mean time.

(a) Compute  $X, Y, Z, I$  and  $\sin x$  for the fixed point  $(\lambda_o, \theta_o)$  and height = 300 km as in Example I, step (d).

(b) 1. Evaluate  $G_k(\lambda_o, \theta_o), a_j(\lambda_o, \theta_o), b_j(\lambda_o, \theta_o)$  as in Example II, steps (c) and (d) for map of foF2 median.

2. Evaluate  $G_k(\lambda_o, \theta_o), a_j(\lambda_o, \theta_o), b_j(\lambda_o, \theta_o)$  as in Example II, steps (c) and (d) for map of M(3000)F2 median.

(c) Compute  $T = 15 h_\ell + 180 - \theta_o$ .

(d) Evaluate  $\cos jT$  and  $\sin jT$ ,  $j = 1, 2, \dots, H$ ,

where  $H$  is the maximum number of harmonics used for either of the characteristics foF2 and M(3000)F2 median.

(e) Compute  $\Omega_f(\lambda_o, \theta_o, T)$  and  $\Omega_m(\lambda_o, \theta_o, T)$  as in Example II, step (g).

(f)  $MUF(4000)F2 = 1.1 \Omega_f \Omega_m$ .

Repeat steps (c) - (f) for each instant of local mean time  $h_\ell$ ; steps (a) and (b) need not be repeated if  $\lambda_o, \theta_o$  remain fixed.

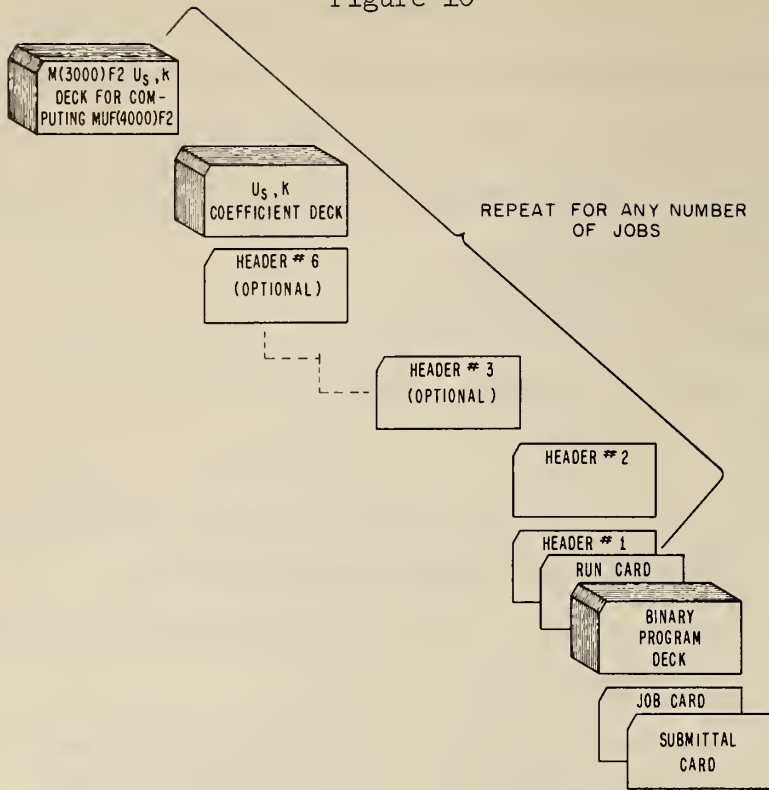
## Sample Program

A sample program has been reproduced here to illustrate the examples of the previous section. This program was written in "Fortran 3600" for the CDC 3600. Program "SAMPLE" illustrates only a few of the possible applications of numerical maps.

The card deck set-up is pictured in Figure 10. Header card 1 is designed to give the following options:

Card	Column	Purpose
1		= 0 prints diurnal curve of desired characteristic for each geographic point (example 2, 3)
		≠ 0 fixed time, compute desired characteristic for each geographic point.
2		= 1 variable characteristic is computed (for example, if the coefficients read in represent foF2, foF2 is printed out)
		= 2 MUF(ZERO)F2 is computed
		= 3 MUF(4000)F2 is computed
3		= 0 in Universal time (in hours)
		≠ 0 Local mean time (in hours)
4 - 6		= number of geographic points punched in header cards 2, 3, ...
7, 8		= number of fixed times (if cc 1 ≠ 0)
9 - 11		= fixed time in format XX.X (if cc 8 = 1)
12 - 14		= fixed time in format XX.X (if cc 8 = 2)
.	.	.
.	.	.
.	.	.
78 - 80		= fixed time in format XX.X (if cc 7, 8 = 24)

Figure 10



Header card number 2 is as follows:

Card Column	Purpose
{ 1 - 4	geographic latitude of first point in format $\pm XX.X$
{ 5 - 8	east geographic longitude of first point in format XXX.X
{ 11 - 14	latitude of 2nd point
{ 15 - 18	longitude of 2nd point
·	·
·	·
·	·
{ 73 - 76	latitude of 20th point
{ 77 - 80	longitude of 20th point

If more than 20 points, punch header card number 3, 4, etc. in the same manner.



A sample printout, corresponding to Example II, is shown in Table 9.

Coefficients used were shown in Table 8.

Table 9

Sample Computer Printout of MUF(ZERO)F2

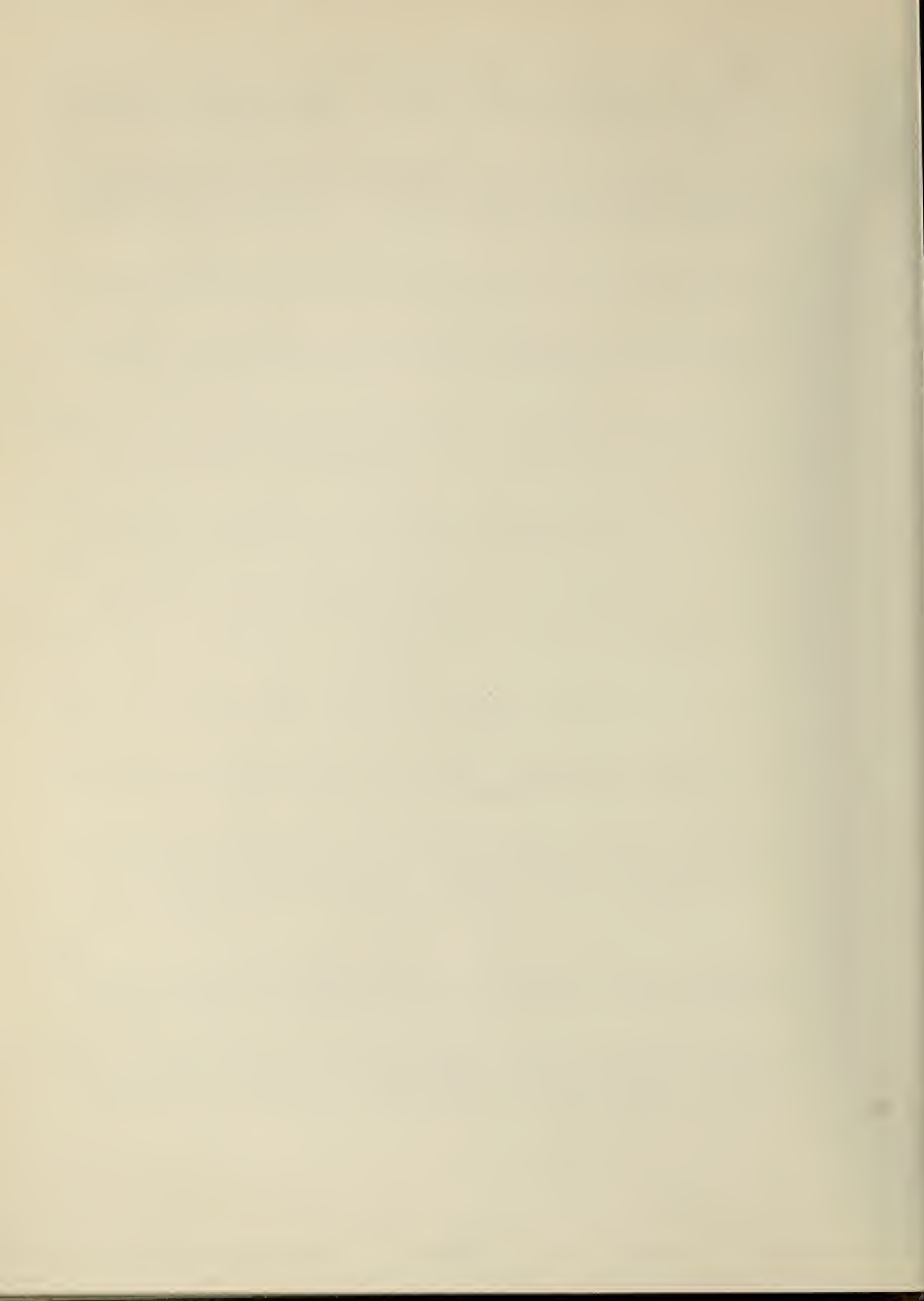
CARD COLUMNS	1,2,3 =	1	2	0	DATE 5803	CHARACTERISTIC	4501
U.T.=	0.0	LATITUDE	86.4	EAST	LONGITUDE	296.6	OMEGA = 6.6
U.T.=	0.0	LATITUDE	76.6	EAST	LONGITUDE	291.3	OMEGA = 6.9
U.T.=	0.0	LATITUDE	41.8	EAST	LONGITUDE	12.5	OMEGA = 7.6
U.T.=	0.0	LATITUDE	-34.5	EAST	LONGITUDE	301.5	OMEGA = 15.5
U.T.=	6.0	LATITUDE	86.4	EAST	LONGITUDE	296.6	OMEGA = 7.0
U.T.=	6.0	LATITUDE	76.6	EAST	LONGITUDE	291.3	OMEGA = 6.7
U.T.=	6.0	LATITUDE	41.8	EAST	LONGITUDE	12.5	OMEGA = 8.8
U.T.=	6.0	LATITUDE	-34.5	EAST	LONGITUDE	301.5	OMEGA = 12.4
U.T.=	12.0	LATITUDE	86.4	EAST	LONGITUDE	296.6	OMEGA = 7.2
U.T.=	12.0	LATITUDE	76.6	EAST	LONGITUDE	291.3	OMEGA = 6.9
U.T.=	12.0	LATITUDE	41.8	EAST	LONGITUDE	12.5	OMEGA = 14.6
U.T.=	12.0	LATITUDE	-34.5	EAST	LONGITUDE	301.5	OMEGA = 13.0
U.T.=	18.0	LATITUDE	86.4	EAST	LONGITUDE	296.6	OMEGA = 6.8
U.T.=	18.0	LATITUDE	76.6	EAST	LONGITUDE	291.3	OMEGA = 7.2
U.T.=	18.0	LATITUDE	41.8	EAST	LONGITUDE	12.5	OMEGA = 11.6
U.T.=	18.0	LATITUDE	-34.5	EAST	LONGITUDE	301.5	OMEGA = 16.0

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Appendix A  
Fourier Coefficients  
for  
foF2 Monthly Median

Figure A - 1

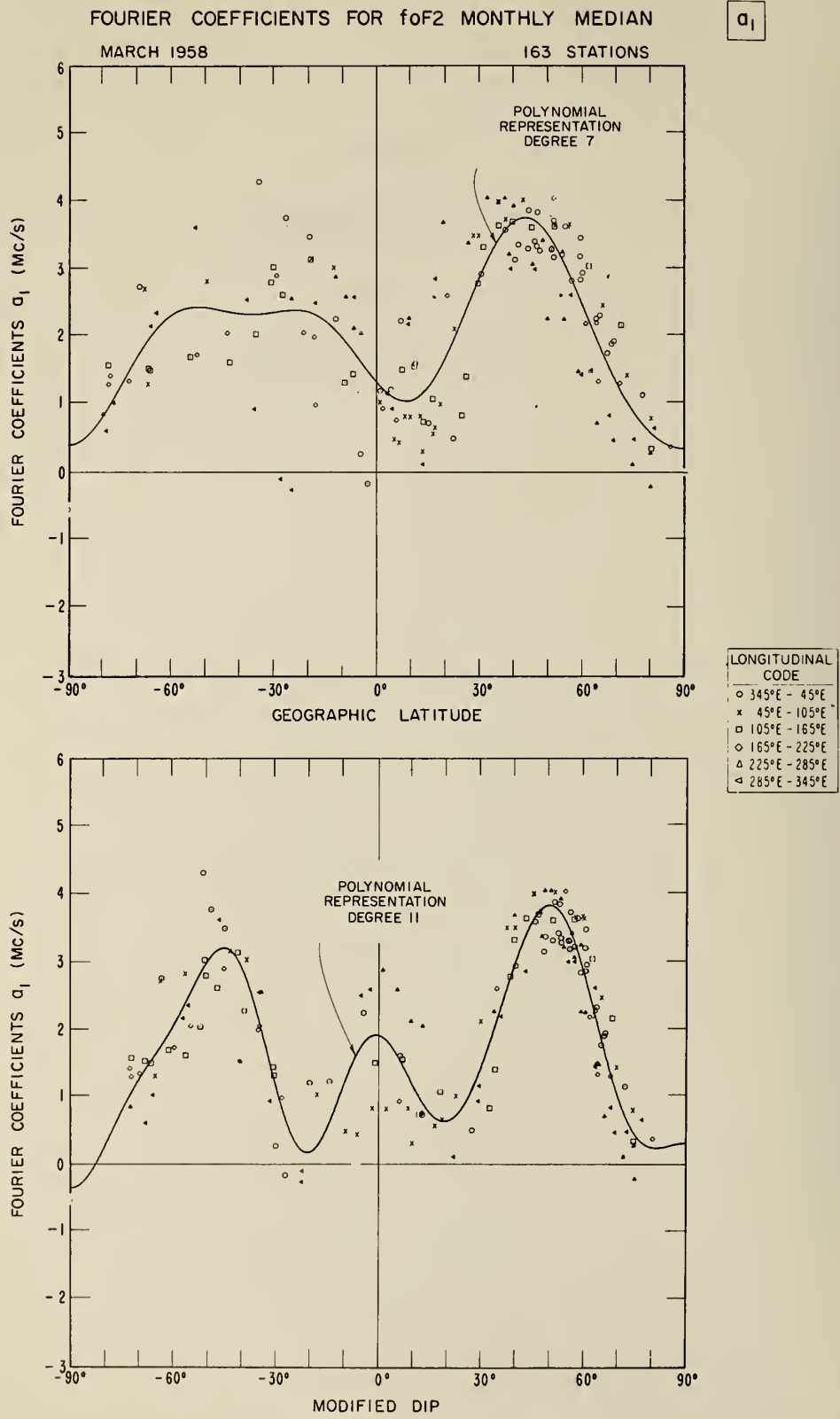


Figure A - 2

FOURIER COEFFICIENTS FOR foF2 MONTHLY MEDIANS  
MARCH 1958 163 STATIONS

$b_1$

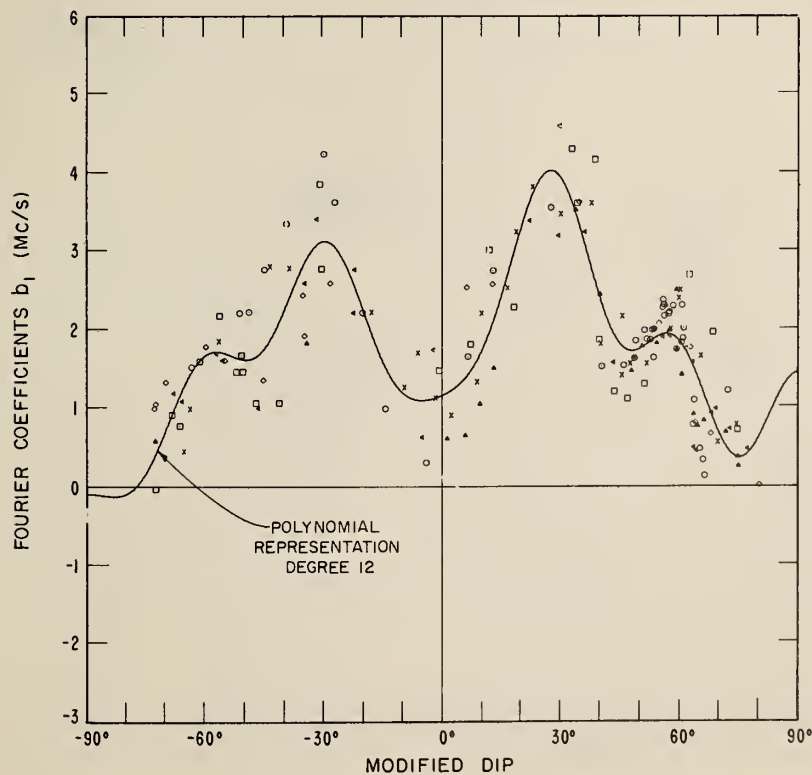
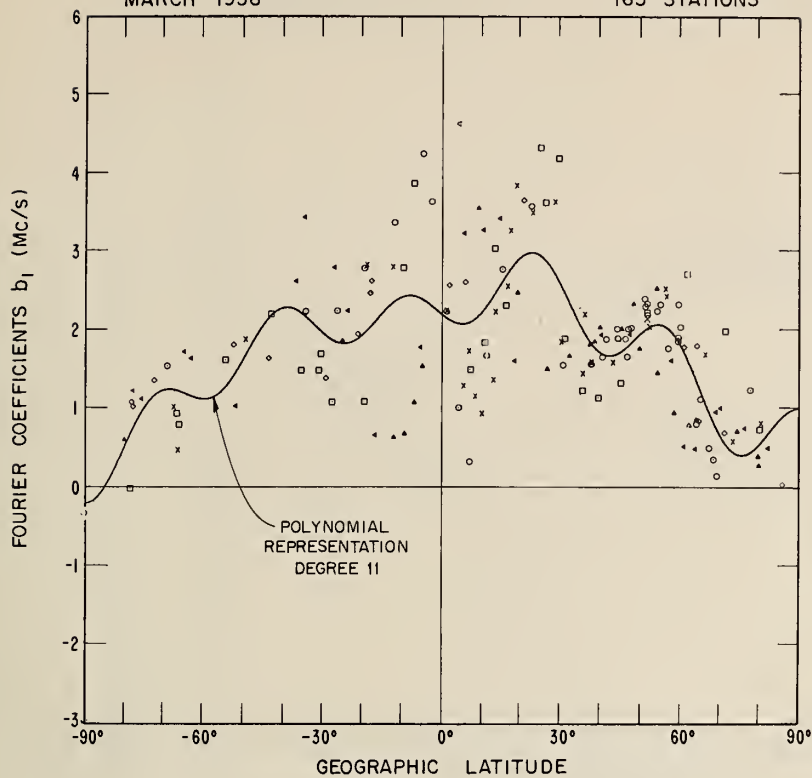
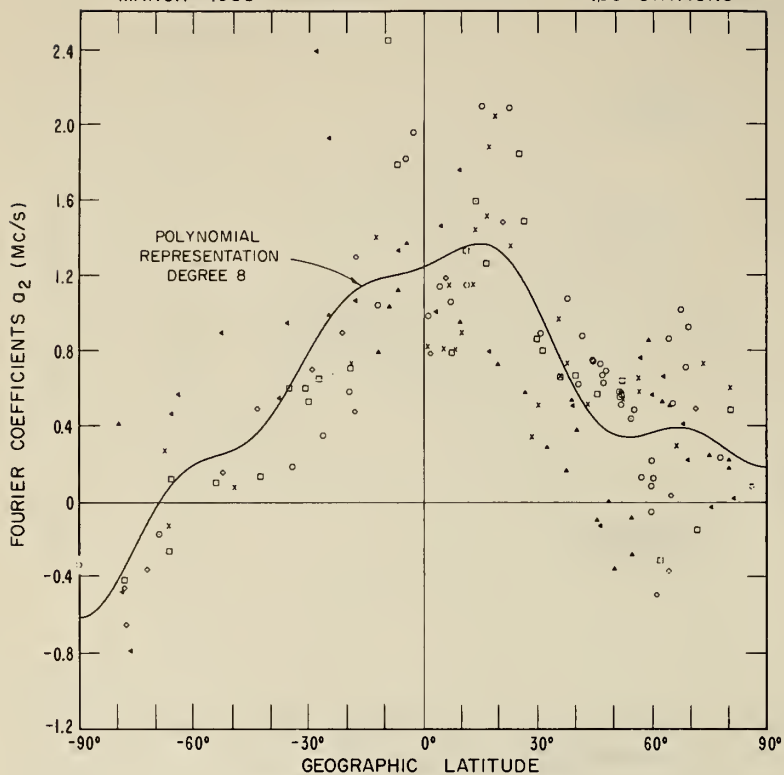


Figure A - 3

FOURIER COEFFICIENTS FOR foF2 MONTHLY MEDIANS  
MARCH 1958 163 STATIONS

$a_2$



LONGITUDINAL CODE	
□	345°E - 45°E
x	45°E - 105°E
○	105°E - 165°E
◇	165°E - 225°E
△	225°E - 285°E
◁	285°E - 345°E

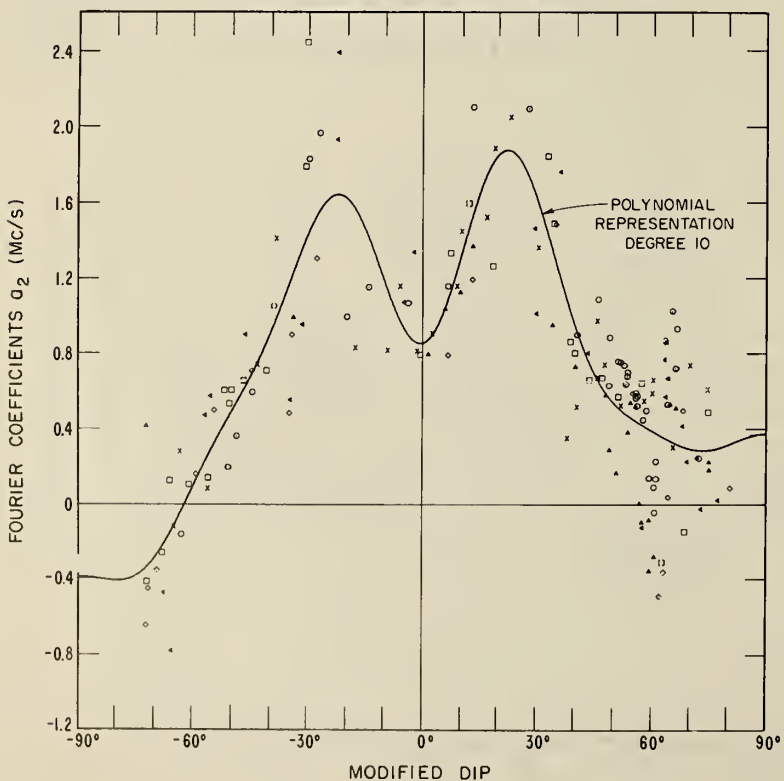
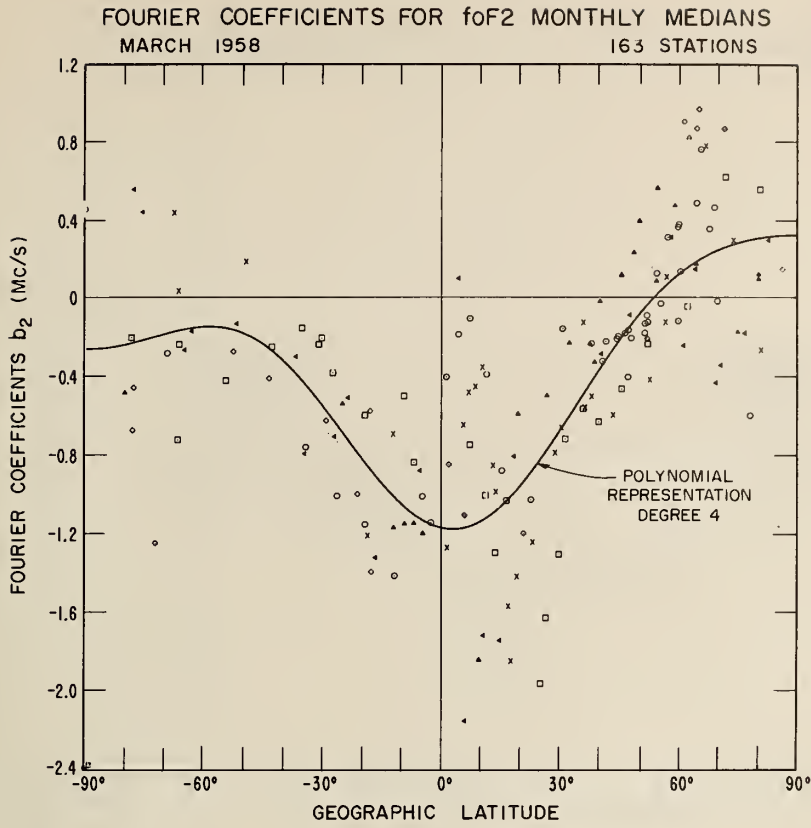




Figure A - 4



$b_2$

LONGITUDINAL CODE	
○	345°E - 45°E
×	45°E - 105°E
□	105°E - 165°E
◇	165°E - 225°E
△	225°E - 285°E
◊	285°E - 345°E

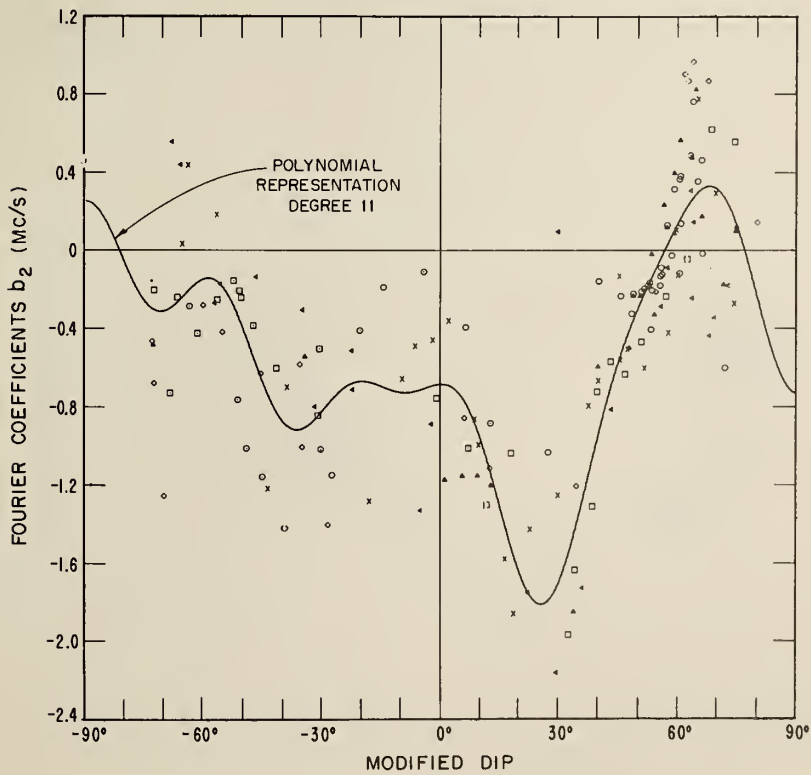


Figure A - 5

FOURIER COEFFICIENTS FOR foF2 MONTHLY MEDIANS  
MARCH 1958 163 STATIONS

$a_3$

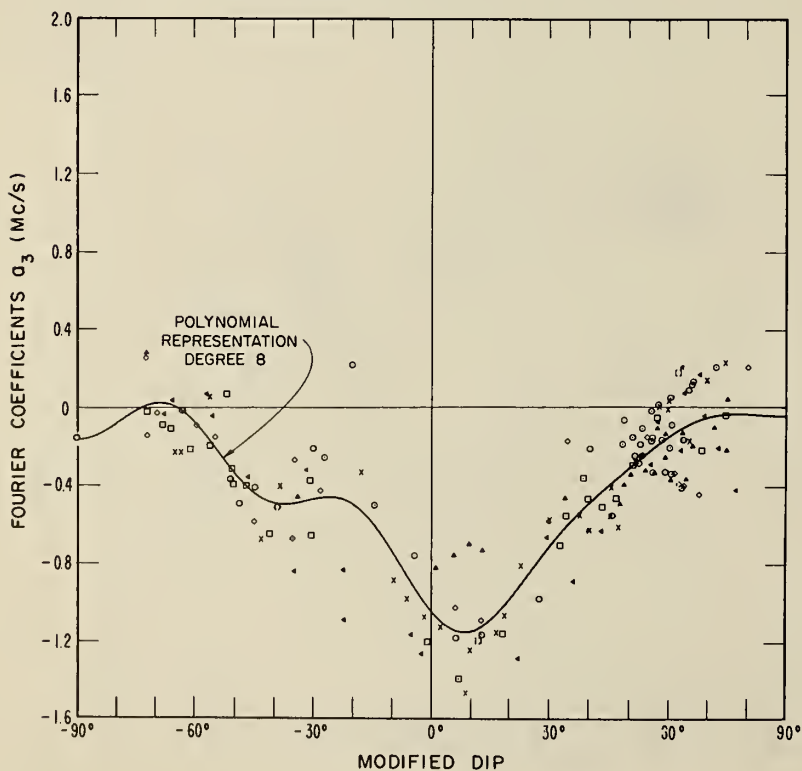
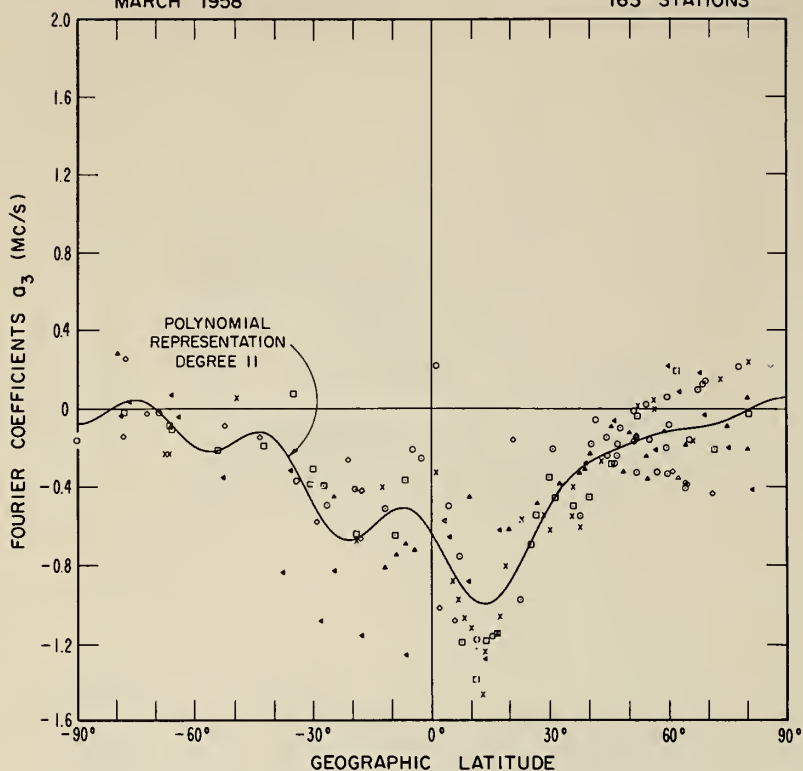
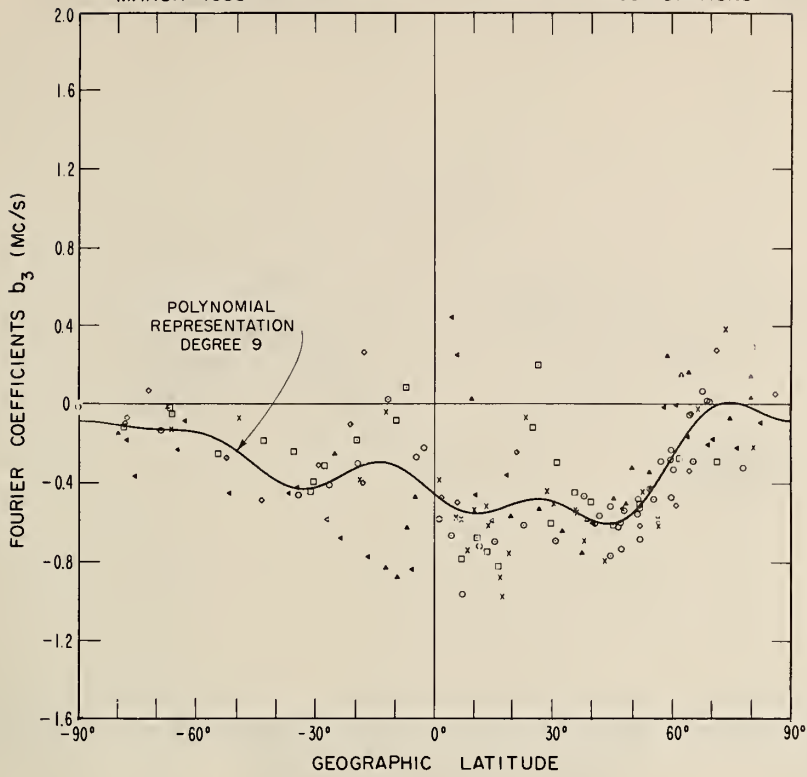


Figure A - 6

FOURIER COEFFICIENTS FOR foF2 MONTHLY MEDIANS  
MARCH 1958 163 STATIONS

b<sub>3</sub>



LONGITUDINAL CODE	
○	345°E - 45°E
×	45°E - 105°E
□	105°E - 165°E
◇	165°E - 225°E
△	225°E - 285°E
◁	285°E - 345°E

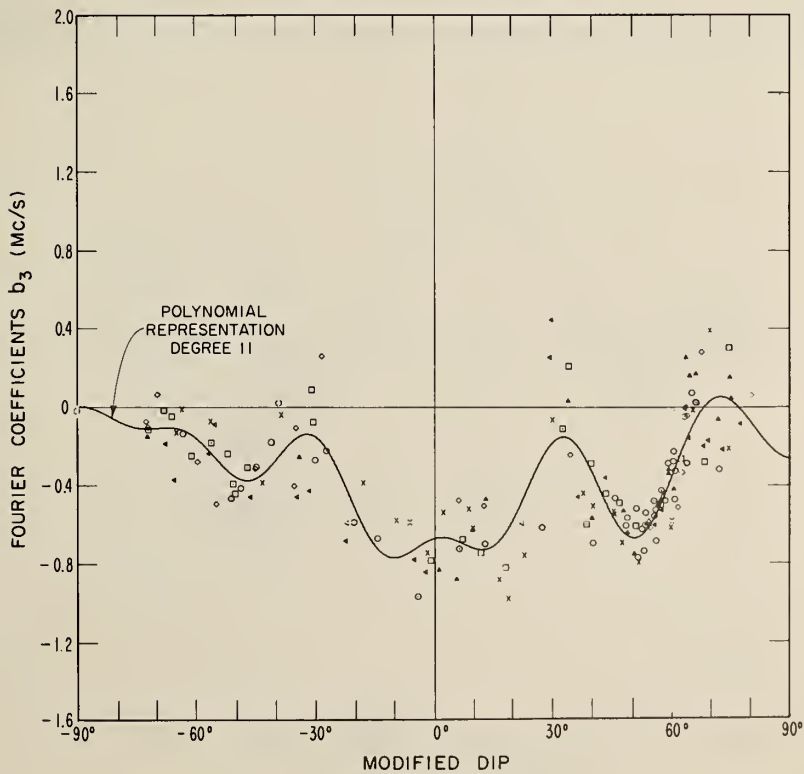


Figure A - 7

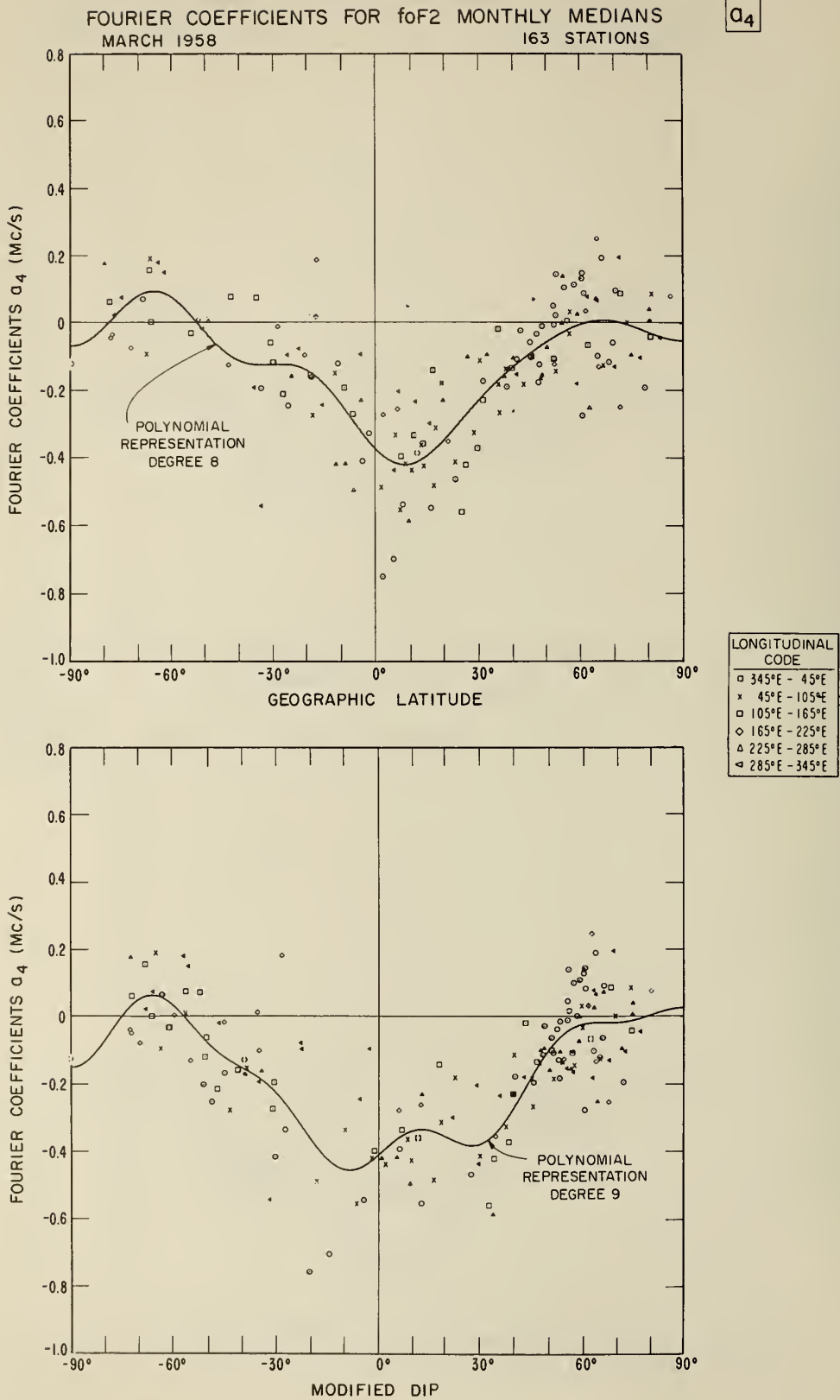
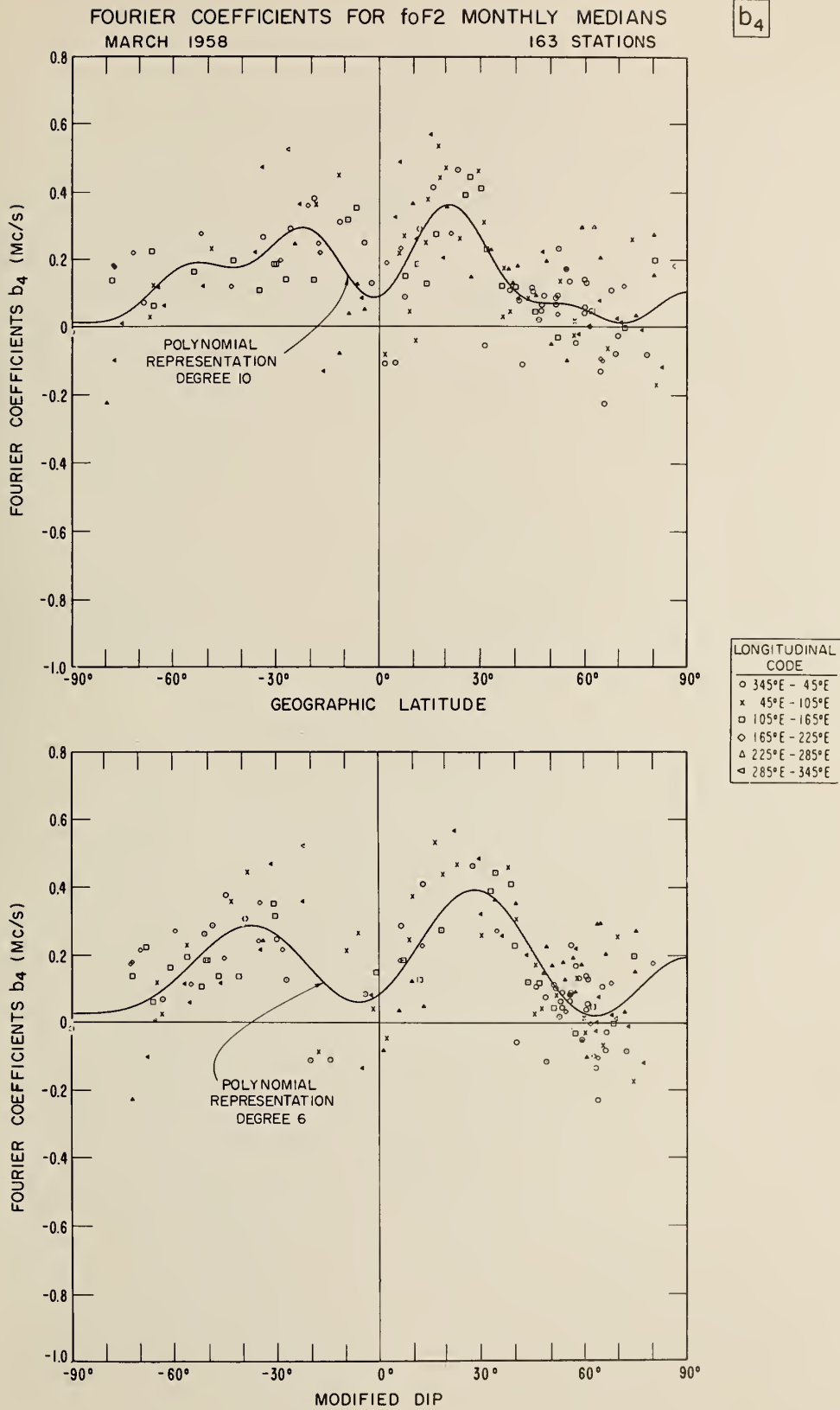
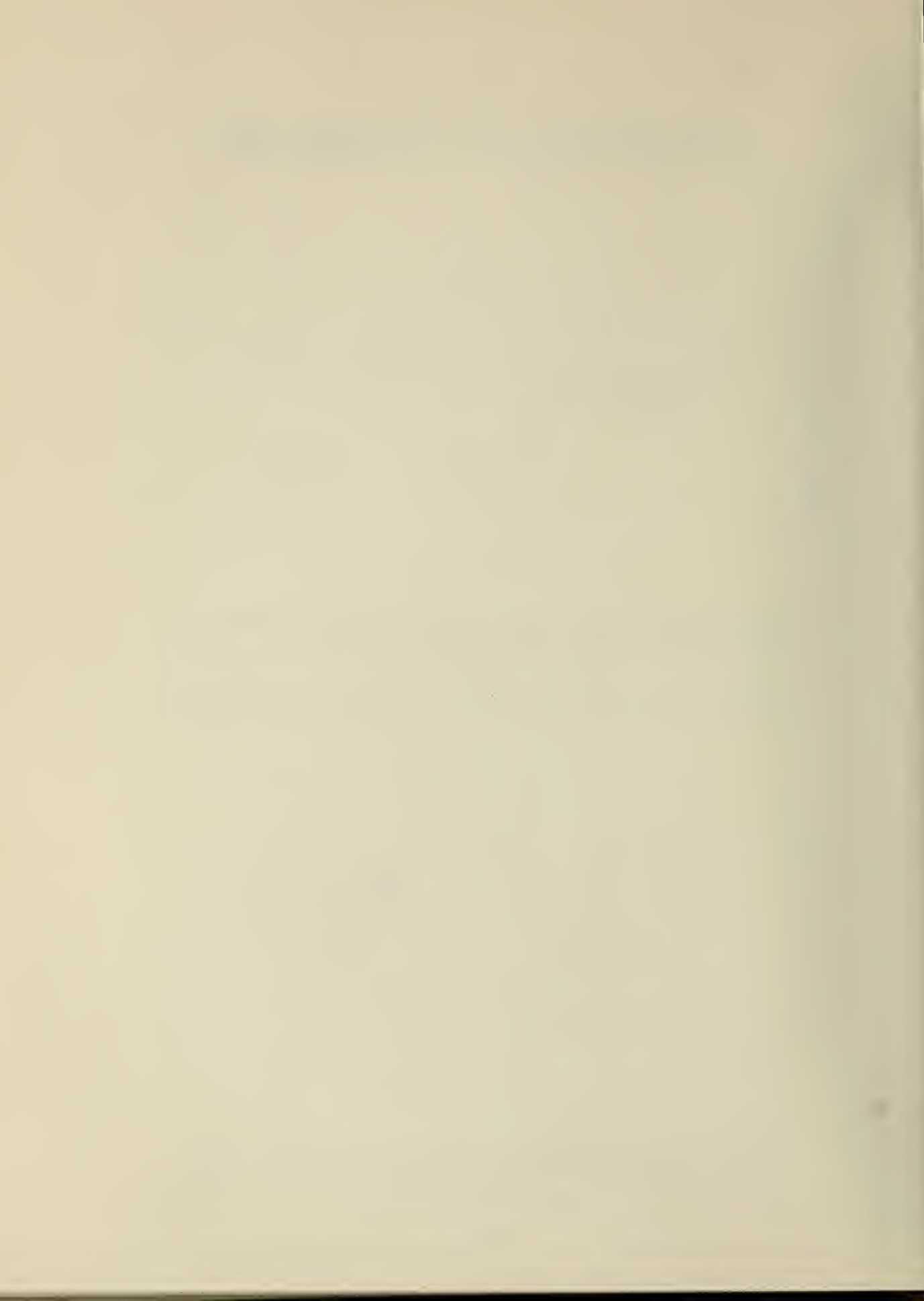


Figure A - 8

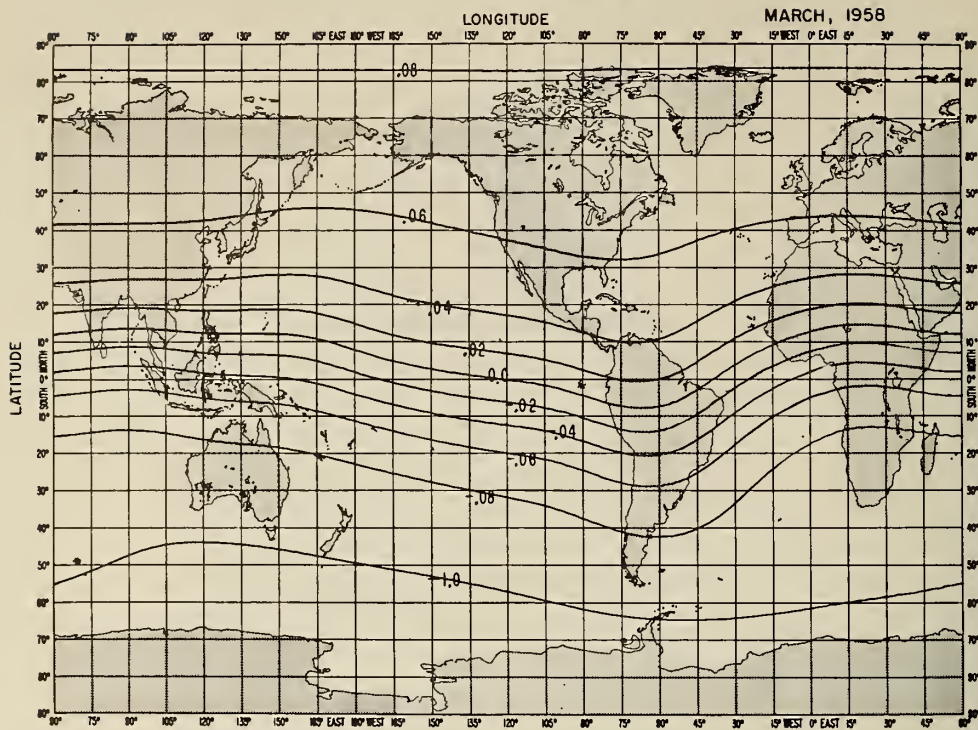




Appendix B  
Contour Maps  
of  
Orthonormal Functions

Figure B - 1

CONTOUR MAP OF ORTHONORMAL FUNCTION  $F_1(\lambda, \theta)$   
(HIGHEST ORDER TERM  $\sin x$ )



CONTOUR MAP OF ORTHONORMAL FUNCTION  $F_2(\lambda, \theta)$   
(HIGHEST ORDER TERM  $\sin^2 x$ )

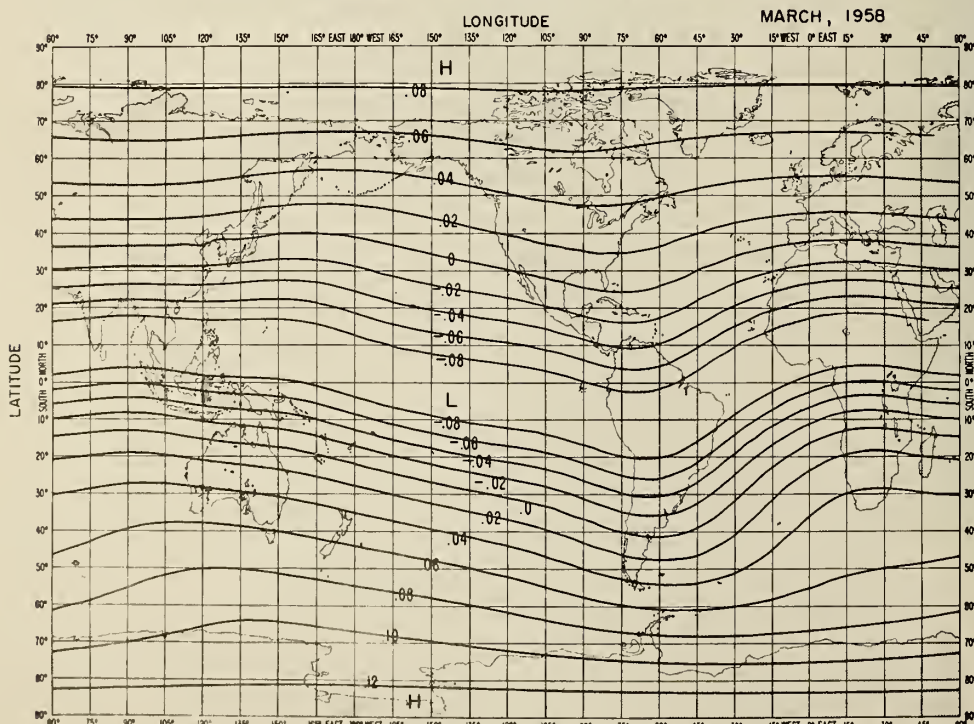
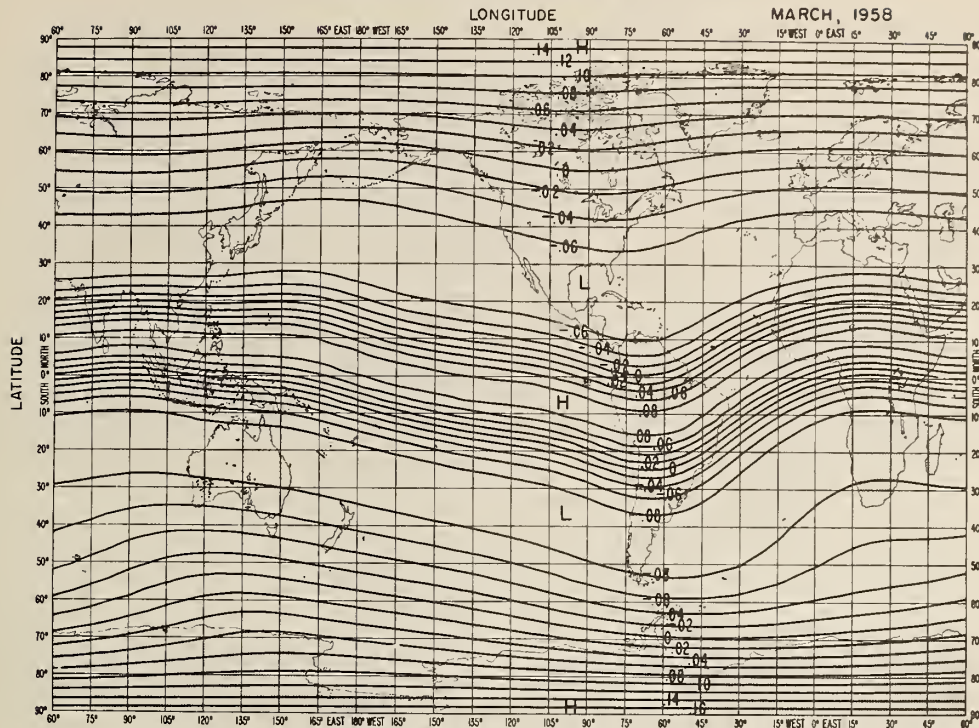




Figure B - 2

CONTOUR MAP OF ORTHONORMAL FUNCTION  $F_4(\lambda, \theta)$   
(HIGHEST ORDER TERM  $\sin^4 x$ )



CONTOUR MAP OF ORTHONORMAL FUNCTION  $F_5(\lambda, \theta)$   
(HIGHEST ORDER TERM  $\sin^5 x$ )

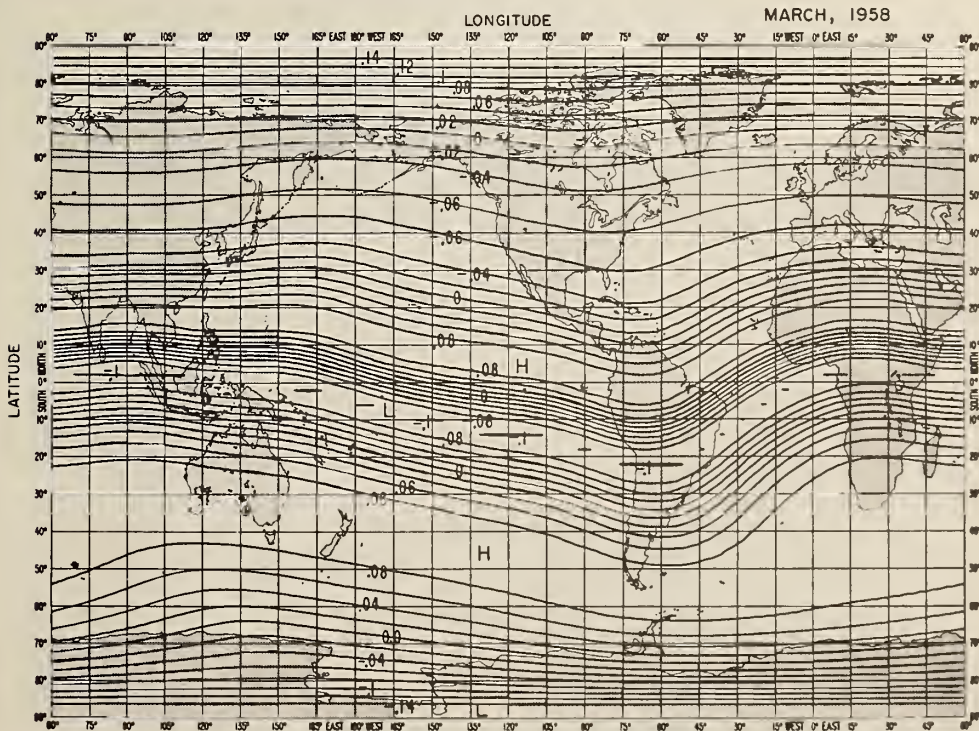
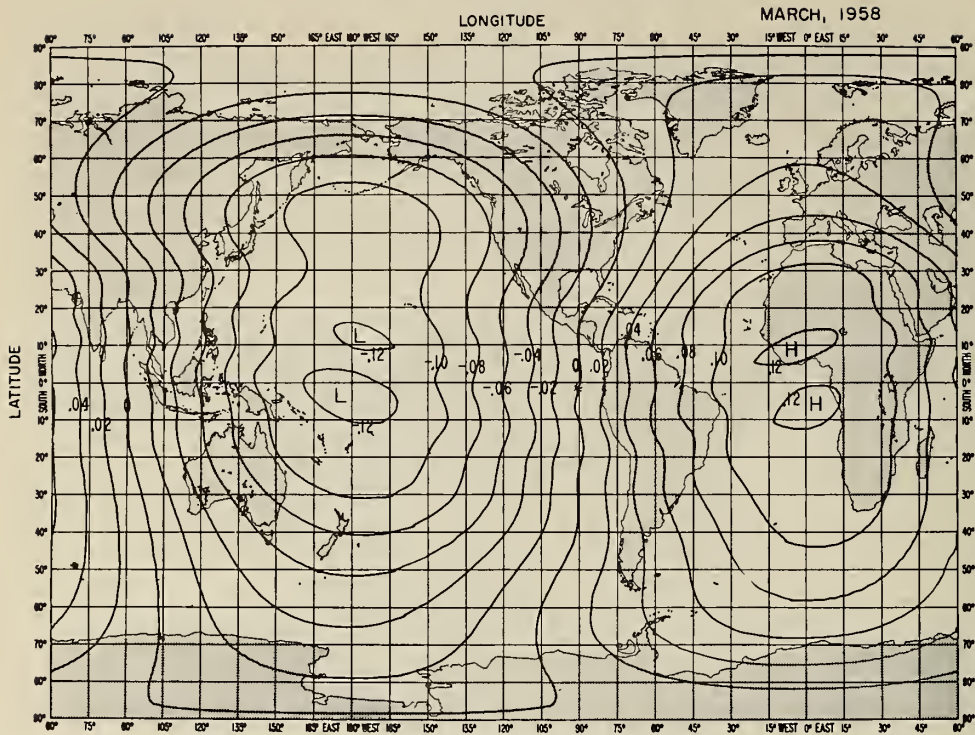


Figure B - 3

CONTOUR MAP OF ORTHONORMAL FUNCTION  $F_{12}(\lambda, \theta)$   
(HIGHEST ORDER TERM  $\cos \lambda \cos \theta$ )



CONTOUR MAP OF ORTHONORMAL FUNCTION  $F_{13}(\lambda, \theta)$   
(HIGHEST ORDER TERM  $\cos \lambda \sin \theta$ )

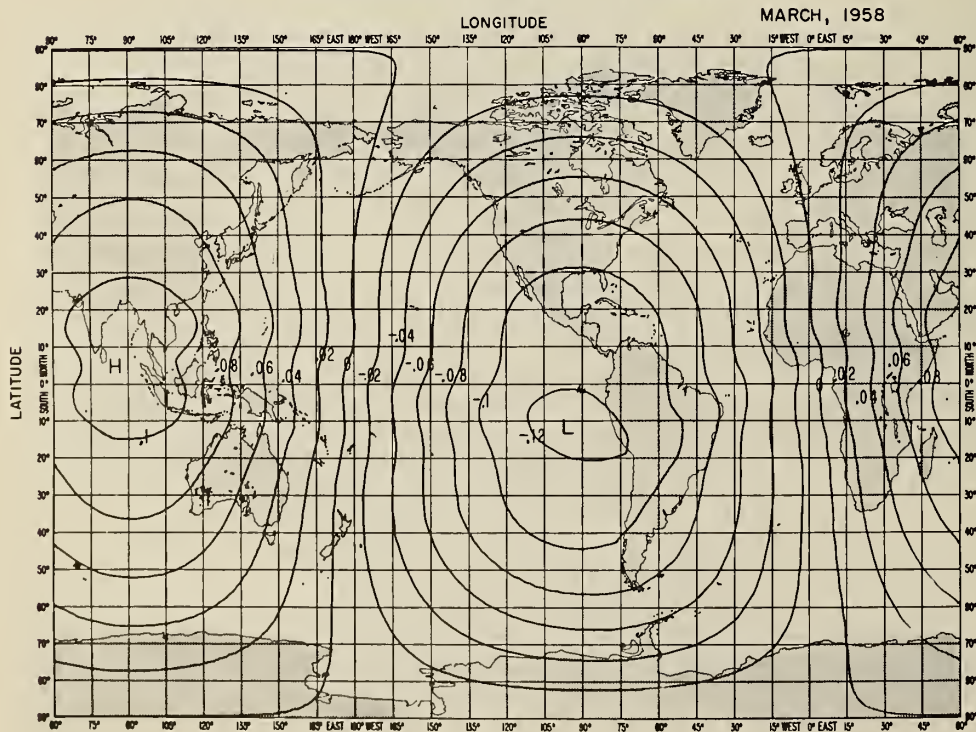
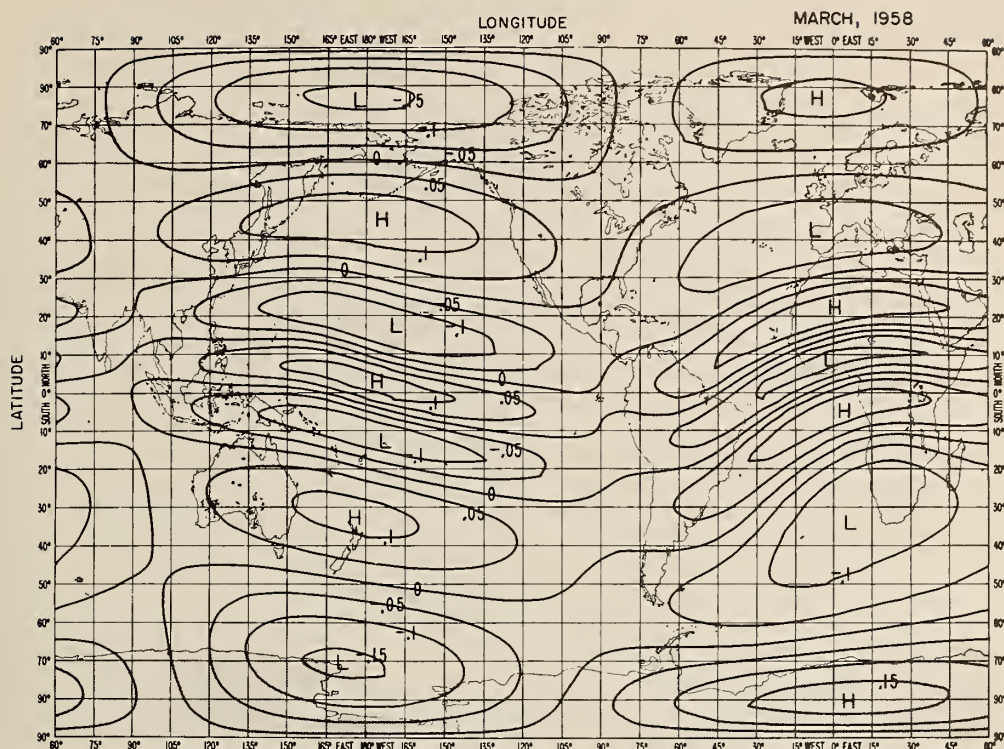
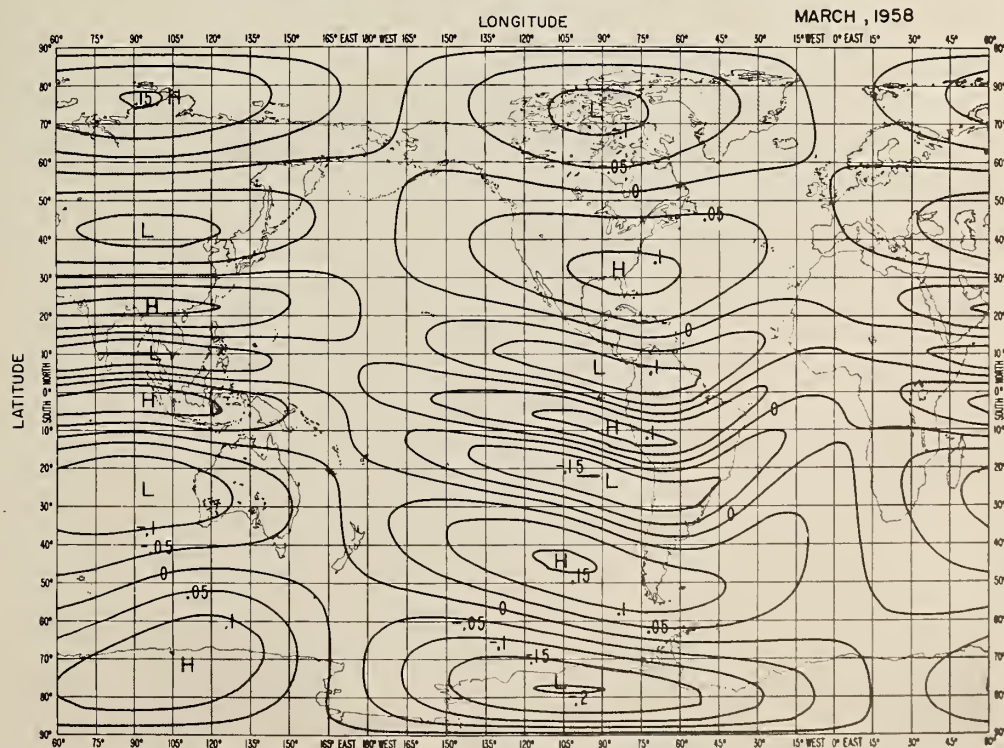


Figure B - 4

CONTOUR MAP OF ORTHONORMAL FUNCTION  $F_{24}(\lambda, \theta)$   
(HIGHEST ORDER TERM  $\sin^6 x \cos \lambda \cos \theta$ )



CONTOUR MAP OF ORTHONORMAL FUNCTION  $F_{25}(\lambda, \theta)$   
(HIGHEST ORDER TERM  $\sin^6 x \cos \lambda \sin \theta$ )





Appendix C  
Diurnal Variation  
of  
Orthonormal Coefficients  $d_k$

Figure C - 1

DIURNAL VARIATION OF ORTHONORMAL COEFFICIENTS  $d_k$

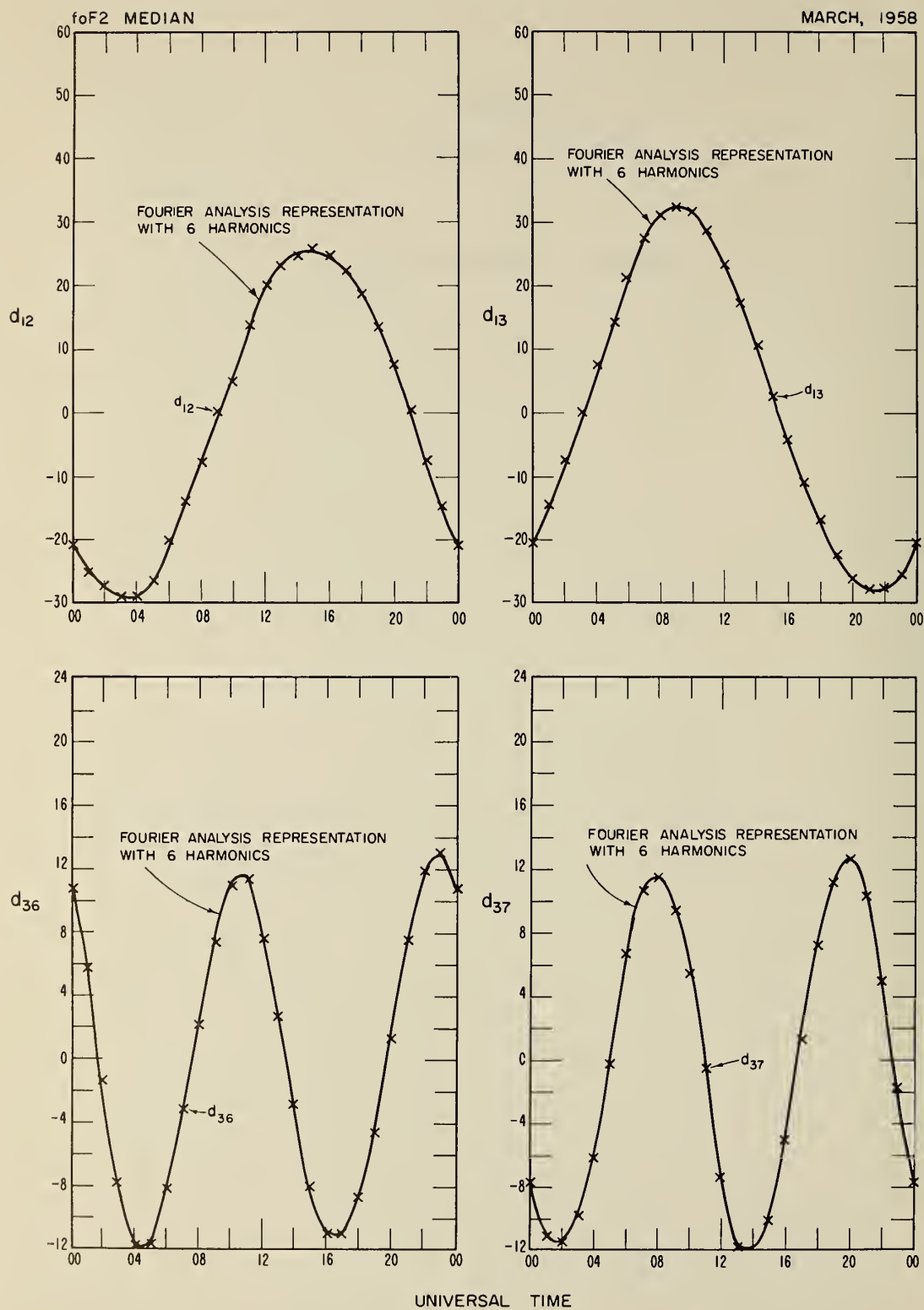
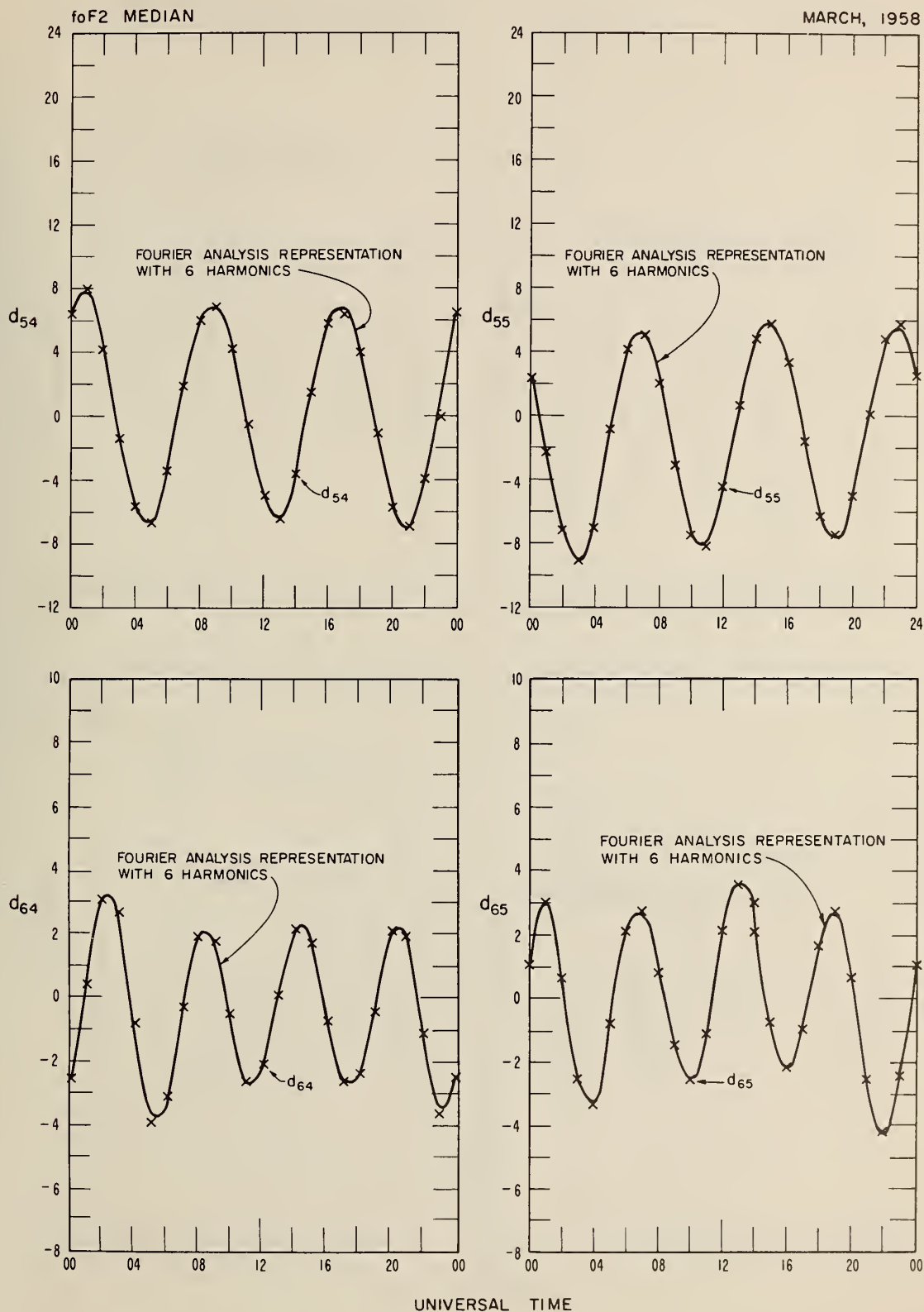


Figure C - 2

DIURNAL VARIATION OF ORTHONORMAL COEFFICIENTS  $d_k$



UNIVERSAL TIME

Figure C - 3

DIURNAL VARIATION OF ORTHONORMAL COEFFICIENTS  $d_k$

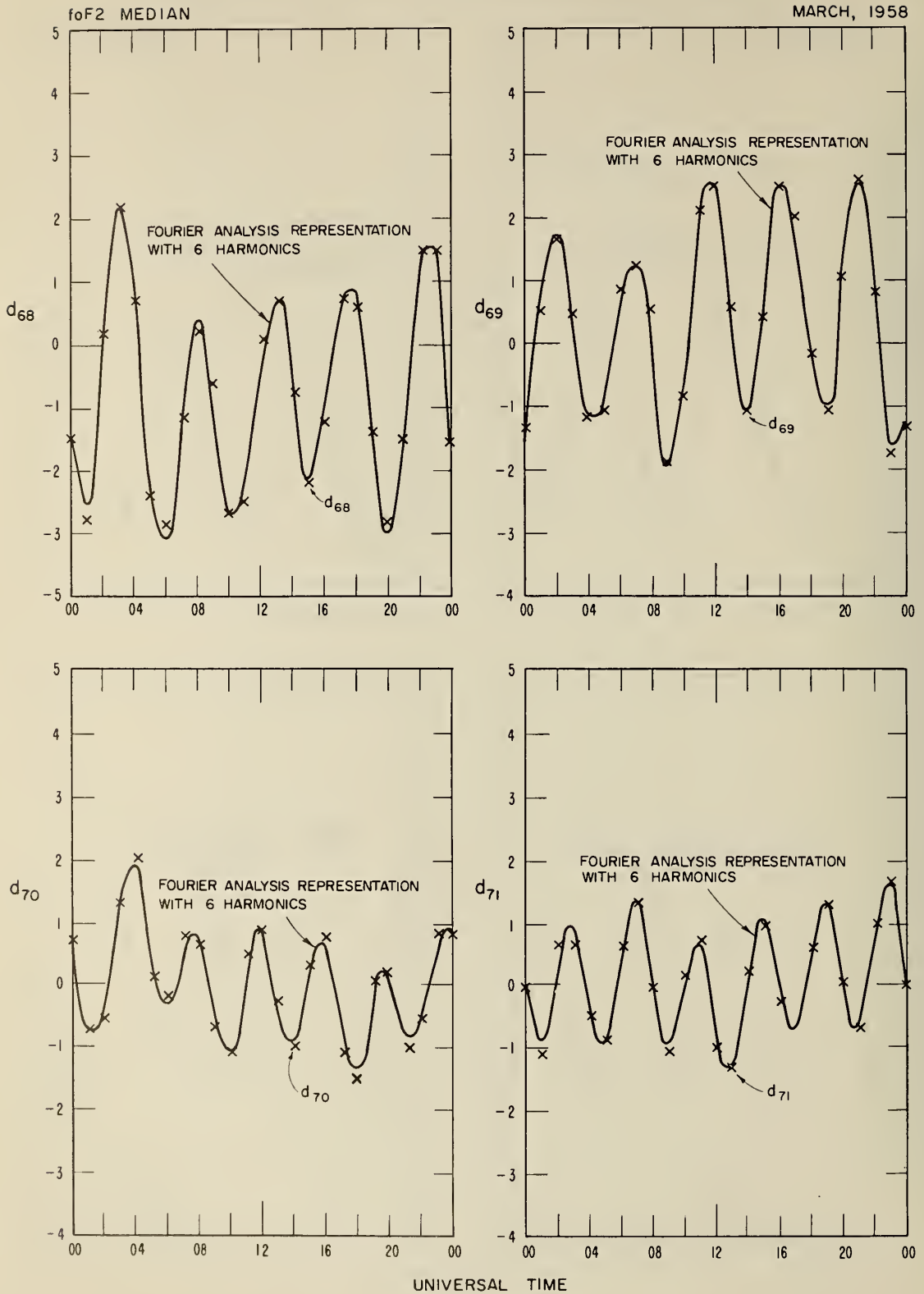
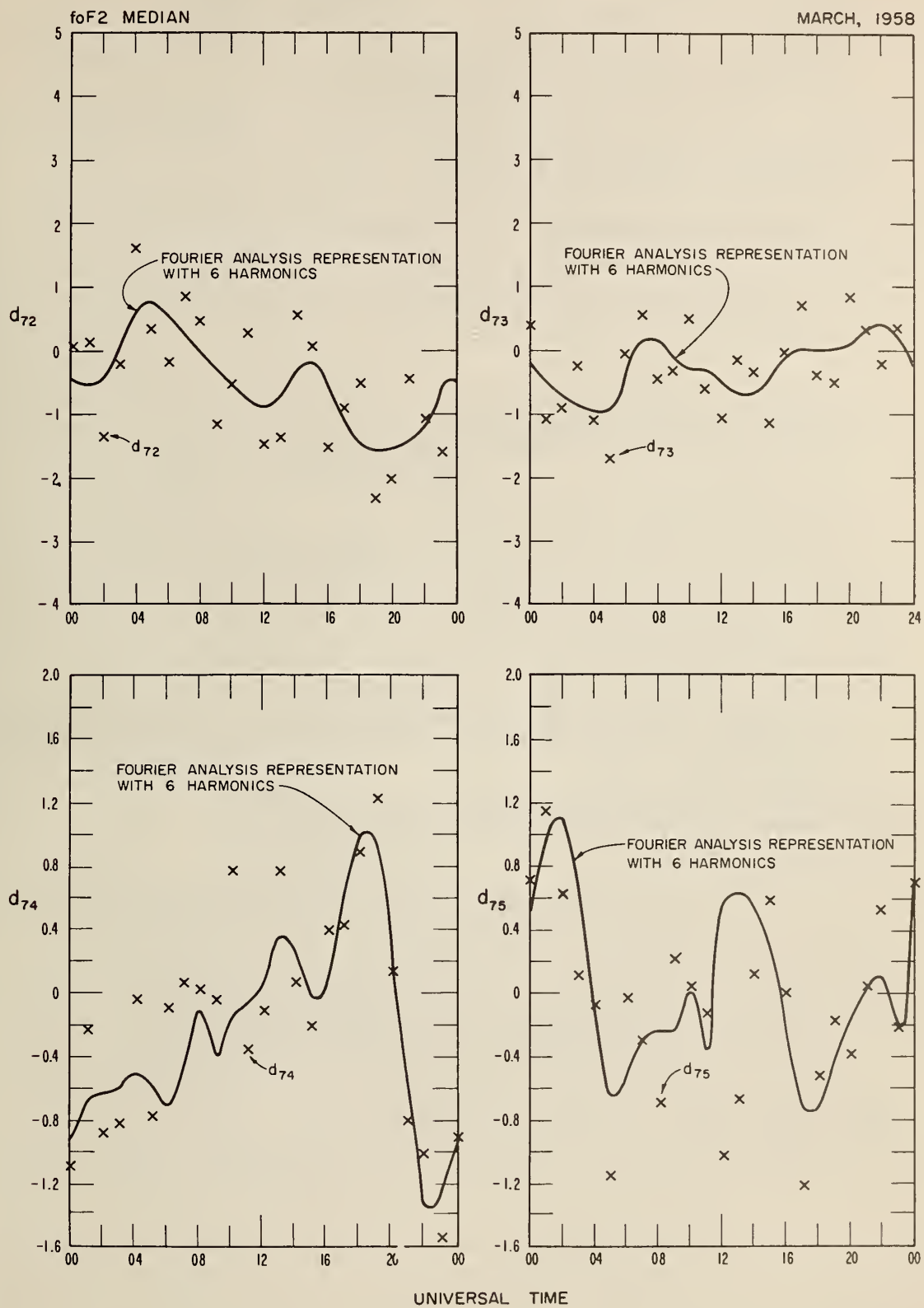




Figure C - 4

DIURNAL VARIATION OF ORTHONORMAL COEFFICIENTS  $d_k$



## Appendix D. SAMPLE Program Listing

```

PROGRAM SAMPLE
TYPE INTEGER CC1,CC2,CC3 $ TYPE REAL M3000
DIMENSION U(17,76),UM(17,76),FLON(100),FLAT(100),K(14),KM(14),
1 COT(8),SIT(8),G(76),GM(76),P(3),C(3),COM(3),AF(9),BF(9),AM(9),
2 BM(9),OMEG (24),TIME(24),DF(76),DM(76)
DATA(P(3) = 3.E5),(DR=.01745329),(RD=57.29578)
C READ OPTIONS CARD
10 READ(60,1) CC1,CC2,CC3,NP,NT,TIME $ IF(EOF,60)12,14
1 FORMAT(3I1,I3,I2,24F3.1)
12 CALL EXIT
C READ POINT CARDS
14 READ(60,2) (FLAT(N),FLON(N),N=1,NP)
2 FORMAT(20F4.1)
C READ COEFFICIENTS
CALL READ U(K,U) $ NFF=K(9)+1 $ GO TO (15,16,18),CC2
15 ICODE=K(12) $ GO TO 20
16 ICODE=4501$ GO TO 20
C READ M(3000)COEFFICIFNTS
18 CALL READ U(KM,UM) $ NMF=KM(9)+1 $ ICODE=4508
20 PRINT 3,CC1,CC2,CC3,K(11),ICODE,K(13),K(14) $ IF(CC1)200,100
3 FORMAT(21H1CARD COLUMNS 1,2,3 =3I3,3X,4HDATEI5,3X,14HCHARACTERISTI
1CI5,3X,13HFUNCTION CODEI4,3X,14HSUNSPOT NUMBERI5/1X)
C CC1=0, COMPUTE DIURNAL CURVE, EACH PT.
100 DO 199 N=1,NP $ P=FLAT(N) $ P(2)=FLON(N) $ CALL MAGFIN(P,COM)
TMP=COM(2)*COM(2)+COM(3)*COM(3)
C=RD*ATANF(ATANF(-COM/SQRTF(TMP))/SQRTF(COSF(DR*P))) $C(2)=FLON(N)
C(3)=P $ CALL GK(K,C,G) $ GO TO(115,110,105),CC2
105 CALL GK(KM,C,GM) $ CALL AJBJ(KM(10),NMF,UM,GM,AM,BM) $ GO TO 115
110 GYF=1.4*SQRTF(COM*COM+TMP)
115 CALL AJBJ(K(10),NFF,U,G,AF,BF) $ IF(CC3)125,120
120 T=-180. $ GO TO 130
125 T=180.--FLON(N)
130 DO 179 I=1,24 $ CALL SICOJT(8,COT,SIT,T)
CALL ABSICO(K(10),AF,BF,COT,SIT,OMEG(I))$GO TO (179,140,135),CC2
135 CALL ABSICC(KM(10),AM,BM,COT,SIT,M3000)$OMEG(I)=1.1*M3000*OMEG(I)
GO TO 179
140 OMEG(I)=OMEG(I)+GYF
179 T=T+15. $ PRINT 4,P(1),P(2)
4 FORMAT( 9H LATITUDEF5.1,3X, 9HLONGITUDEF6.1)
199 PRINT 5,OMEG $ GO TO 10
5 FORMAT(3X,24F5.1)
C CC1 NOT=0, COMPUTE OMEGA FOR FIXED TIMES, VARYING POINTS
200 DO 279 I=1,NT $ T=15.*TIME(I)-180. $ CALL SICOJT(8,COT,SIT,T)
CALL DKSICO(NFF,K(10),U,SIT,COT,DF) $ IF(CC2.EQ.3)205,210
205 CALL DKSICO(NMF,KM(10),UM,SIT,COT,DM)
210 DO 279 N=1,NP $ P=FLAT(N) $ P(2)=FLON(N) $ CALL MAGFIN(P,COM)
TMP=COM(2)*COM(2)+COM(3)*COM(3) $ C(2)=FLON(N) $ C(3)=P
C=RD*ATANF(ATANF(-COM/SQRTF(TMP))/SQRTF(COSF(DR*P)))
CALL GK(K,C,G) $CALL DKGK(NFF,G,DF,OMEGA) $GO TO(279,220,215),CC2
215 CALL GK(KM,C,GM) $ CALL DKGK (NMF,GM,DM,M3000)
OMEGA=1.1*M3000*OMEGA $ GO TO 279
220 OMEGA=OMEGA+1.4*SQRTF(COM*COM+TMP)
279 PRINT 6,TIME(I),P(1),P(2),OMEGA $ GO TO 10
6 FORMAT(6H U.T.=F5.1,3X, 8HLATITUDEF5.1,3X,14HEAST LONGITUDEF6.1,3X

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1 ,7HOMEGA =F6.1)
  END
  SUBROUTINE MAGFIN(POS,UNE)
    COMPUTE NASA MAGNETIC FIELD COMPONENTS
C   DIMENSION P(7,7),DP(7,7),CP(7),AOR(7),SP(7),POS(3),UNE(3),CT(7,7),
1   G(7,7),H(7,7)
  DATA (P(1,1)=1.), (DP(1,1)=0.), (SP(1)=0.), (CP(1)=1.), (CT=2(0.)),
1.333333333, .266666667, .25714286, .25396825, .25252525, 3(0.), .2,
2.22857142, .23809523, .24242424, 4(0.), .14285714, .19047619, .21212121,
35(0.), .11111111, .16161616, 6(0.), .09090909, 14(0.), (G=0., .304112,
4.024035, -.031518, -.041794, .016256, -.019523, 0., .021474, -.051253,
5.062130, -.045298, -.034407, -.004853, 2(0.), -.013381, -.024898,
6-.021795, -.019447, .003212, 3(0.), -.006496, .007008, -.000608, .021413,
74(0.), -.002044, .002775, .001051, 5(0.), .000697, .000227, 6(0.), .001115
8), (H=8(0.), -.057989, .033124, .014870, -.011825, -.000796, -.005758,
92(0.), -.001579, -.004075, .010006, -.002, -.008735, 3(0.), .00021, .00043
1), .004597, -.003406, 4(0.), .001385, .002421, -.000118, 5(0.), -.001218,
2-.001116, 6(0.), -.000325), (RD=57.2957795), (HC=6371200.)
  P2 = POS(2) $ P1=POS $ IF(P1-89.9)2,4,1
1 P1=89.9 $ P2 = 0 $ GO TO 4
2 IF(P1+89.9)3,4,4
3 P1=-89.9 $ P2 = 0
4 PHI= P2 /RD$AR=HC/(HC+POS(3))$C=SINF( P1 /RD)$S=SQRTF(CP-C*C)
  SP(2)=SINF(PHI)$CP(2)=COSF(PHI)$AOR=AR*AR$AOR(2)=AOR*AR$DO 5 M=3,7
  SP(M)=SP(2)*CP(M-1)+CP(2)*SP(M-1)
  CP(M)=CP(2)*CP(M-1)-SP(2)*SP(M-1)
5  AOR(M)=AR*AOR(M-1)
  BV=0.$BN=BV$BPHI=BV$DO 6 N=2,7$FN=N$SUMR=0.$SUMT=SUMR$SUMP=SUMT
  DO 7 M=1,N $ IF(N-M) 8,9,8
9  P(N,N)=S*P(N-1,N-1)$DP(N,N)=S*DP(N-1,N-1)+C*P(N-1,N-1) $ GO TO 10
8  P(N,M) = C*P(N-1,M) - CT(N,M)*P(N-2,M)
  DP(N,M) = C*DP(N-1,M) - S*P(N-1,M) - CT(N,M)*DP(N-2,M)
10 FM=M-1$TS=G(N,M)*CP(M)+H(N,M)*SP(M)$SUMR=SUMR+P(N,M)*TS
  SUMT=SUMT+DP(N,M)*TS
7  SUMP=SUMP+FM*P(N,M)*(-G(N,M)*SP(M)+H(N,M)*CP(M))
  BV=BV+AOR(N)*FN*SUMR $ BN=BN-AOR(N)*SUMT
6  BPHI=BPHI-AOR(N)*SUMP$UNE(1)=-BV$UNE(2)=BN$UNE(3)=-BPHI/S
  RETURN $ END
  SUBROUTINE GK ( K,C,G )
    COMPUTE COORDINATE FUNCTIONS,G(I),I=1,...,K+1
C   C(1)=MODIFIED LATITUDE,C(2),C(3)=GEOG.LAT.,LONGITUDE,RESPECTIVELY
C   DIMENSION K(1),C(1),G(1)
  DATA (DR=.017453293), (N=8)
  X=DR*C $Y=C(2)*DR$Z=DR*C(3)$K0=K$SX=SINF(X)$G(2)=SX$G=1.
  IF(K0.EQ.1) 12,4
4  DO 10 I=2,K0
10 G(I+1)=SX*G(I)
12 KDIF=K(2)-K0$ IF(KDIF)16,80
16 J=1 $ CX1=COSF(Z) $ CX=CX1 $ T=Y
18 KC=K(J)+4$G(KC-2)=CX*COSF(T) $ G(KC-1)=CX*SINF(T) $IF(KDIF-2)20,28
20 KN=K(J+1) $ DO 22 I=KC,KN,2 $ G(I)=SX*G(I-2)
22 G(I+1)=SX*G(I-1)
28 IF(J.EQ.N)80,30
30 KDIF=K(J+2)-KN $ IF(KDIF)32,80

```

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32 CX=CX*CX1$ J=J+1 $ FJ=J $ T=FJ*Y $ GO TO 18
80 RETURN
END
SUBROUTINE READ U(K,U)
C      READ U(S,K) COEFFICIENTS
      DIMENSION K(1),U(17,1)
      READ (60,1) KD,(K(I),I=1,10),(K(I),I=12,14),DATA,K(11),NCARDS
1  FORMAT(I1,I3I5,6X,A1,I4,I3)
      MX=K(9)+1 $ DO 5 I=1,17 $ DO 5 J=1,MX
5  U(I,J)=0 $ DO 25 I=1,NCARDS
      READ (60,2) I1,M1,D1,I2,M2,D2,I3,M3,D3,I4,M4,D4,DATUM,KDATE,ID
2  FORMAT(4(2I2, E14.7),A1,I4,I3)
      L1=I1+1 $ L2=I2+1 $ L3=I3+1 $ L4=I4+1 $ M1=M1+1 $ M2=M2+1
      M3=M3+1 $ M4=M4+1 $ IF(K(11).EQ.KDATE)10,30
10 IF( DATA.EQ. DATUM)15,30
15 IF(ID.EQ.I)16,30
16 U(L1,M1)=D1 $ IF(I.LT.NCARDS)18,20
18 U(L2,M2)=D2 $ U(L3,M3)=D3 $ GO TO 25
20 IF(I2)21,28
21 U(L2,M2)=D2 $ IF(I3)22,28
22 U(L3,M3)=D3 $ IF(I4)25,28
25 U(L4,M4) = D4
28 RETURN
30 PRINT 3,ID,K(11),K(12)
3  FORMAT(23H1 ERROR IN CARD NUMBER I3,2I5)
      CALL EXIT
      END
SUBROUTINE DKSICO (MX,LH,D,SITIME,COTIME,DK)
C      COMPUTE D SUB K, COEFFICIENTS FOR A FIXED TIME
      DIMENSION D(17, 1),COTIME(1),SITIME(1),DK(1)
      DO 5 K = 1,MX
      DK(K)=D(1,K)
      DO 5 L=1,LH
5  DK(K) = DK(K) + D(2*L,K)*SITIME(L)+D(2*L+1,K)*COTIME(L)
      RETURN
      END
SUBROUTINE DKGK(MX,G,DKSTAR,OMEGA)
C      COMPUTE OMEGA , SUMMING THE GEOGRAPHIC SERIES
      DIMENSION G(1),DKSTAR(1)
      OMEGA = G*DKSTAR
      DO 5 K=2,MX
5  OMEGA = OMEGA+DKSTAR(K)*G(K)
      RETURN
      END
SUBROUTINE AJBJ(LH,Mx,D,G,ASTAR,BSTAR)
C      COMPUTE COEFFICIENTS FOR A FIXED GEOGRAPHIC POINT
      DIMENSION D(17, 1),G( 1),ASTAR(1),BSTAR(1)
      N=LH+1
      DO 5 J=1,N
      ASTAR(J)=0.
      DO 5 K=1,MX
5  ASTAR(J)=ASTAR(J)+D(2*J-1,K)*G(K)
      DO 10 J=2,N
      BSTAR(J)=0.

```

```

DO 10 K=1,MX
10 BSTAR(J)=BSTAR(J)+D(2*J-2,K)*G(K)
RETURN
END
SUBROUTINE ABSICO(LH,ASTAR,BSTAR,COTIME,SITIME,OMEGA)
C   COMPUTE OMEGA , SUMMING THE FOURIER SERIES
DIMENSION COTIME(1),SITIME(1),ASTAR(1),BSTAR(1)
OMEGA=ASTAR
DO 5 J=1,LH
5 OMEGA=OMEGA+ASTAR(J+1)*COTIME(J)+BSTAR(J+1)*SITIME(J)
RETURN
END
SUBROUTINE SICOJT(L,C,S,A)
C   COMPUTE SIN(JT),COS(JT),J=1,...,L FOR ANGLE A
DIMENSION C(1),S(1)
T=.01745329*A $ C=COSF(T) $ S=SINF(T) $ DO 10 I=2,L
C(I)=C*C(I-1)-S*S(I-1)
10 S(I)=C*S(I-1)+S*C(I-1) $ RETURN
END

```





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