JOURNAL OF RESEARCH of the National Bureau of Standards - A. Physics and Chemistry Vol. 72A, No. 4, July-August 1968

# **Synthesis of Cerite**

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#### (April 29, 1968)

Some compounds isostructural with cerite have been synthesized hydrothermally at temperatures from 500 to 720 °C. The compositional range for cerite, and its phase relation to the silicate apatites are established in terms of ionic radii of rare-earth and divalent ions. The proposed formula for cerite on the basis of isomorphism with whitlockite,  $Ca_3(PO_4)_2$ , is discussed. A solid-solution series between  $Ca_9P_6O_{24}-Ca_7Mg_2P_6O_{24}$  (whitlockite - Mg-whitlockite) has been established.

Key Words: Ca, Mg phosphate, hydrothermal synthesis, inorganic synthesis, mineral chemistry, rare-earth silicate, x-ray powder analysis.

## 1. Introduction

During a study of silicate apatites, a very wide compositional range for the cerite phase was found under hydrothermal conditions [1].<sup>1</sup> This range adjoins that for the rare-earth silicate apatites. Despite its early description, the rare-mineral cerite  $[2-7]$ , from which the element cerium was discovered, has not yet been fully characterized, because of the complexity of its composition. In this paper, therefore, I report the conditions for its formation and its compositional range. Recently Keppler [8] reported an isostructural relation between cerite and whitlockite. I suspect that the maximum number of Mg ions in Mg-whitlockite may equal the number of divalent ions in cerite, consequently, the whitlockite-Mgwhitlockite solid-solution series was also investigated and the results are given here.

### 2. Results and Discussion

Cerite was synthesized from hydrosilicates which were precipitated in the ratio  $\dot{M}^{2+}$ :  $M^{3+}$  (i.e., rare earth):  $M^{4+}$  (i.e., Si) = 2:7:6. The precipitates were centrifuged, washed, dried, and in some experiments admixed with carbonates. The synthesis was completed by hydrothermal heating at 2 kbar and 550 to 680  $^{\circ}$ C. Whitlockite was synthesized from stoichiometric mixtures, prepared by the slow evaporation of alkaline solution containing calcium, magnesium nitrate and ammonium dihydrogen phosphate and heated in air at 1050 to 1120 °C. Experimental results are given in table 1 with the unit-cell dimensions calculated accord-

TABLE 1. Results of the synthesis of cerite, whitlockite and Mgwhitlockite and the unit cell dimensions

Formula (tentative)	Temp °C	Pressure	Hour	$\overline{a}$	$\overline{c}$
$Mg_2La_7Si_6O_{23}(OH)_{3}$	550	$\overline{2}$	20	10.77	38.40
$M_{\mathfrak{D}}Nd_{7}Si_{6}O_{23}(OH)_{3}$	700	$\overline{2}$	20	10.64	37.63
$Mg_2Sm_7Si_6O_{23}(OH)_3$	550	$\overline{2}$	20	10.62	36.8
$Mg_2Gd_7Si_6O_{23}(OH)_{3}$	550	$\overline{2}$	20	10.49	36.64
$Mg_2Dy_7Si_6O_{23}(OH)_{3}.\dots$	700	$\overline{2}$	20	10.47	36.32
	680	$\overline{2}$	48	10.44	35.92
	650	$\overline{2}$	20	cerite <sup>a</sup>	
$Ni2Sm7Si6O23(OH)3$	550	$\overline{2}$	20	10.58	36.74
$Ni2H07Si6O23(OH)3$	680	$\overline{2}$	20	10.47	36.34
$Co_2La_7Si_6O_{23}(OH)_3$	550	$\overline{2}$	48	10.78	38.40
	680	$\overline{2}$	20	cerite <sup>a</sup>	
$Co_2Sm_7Si_6O_{23}(OH)_3$	550	$\overline{2}$	20	10.6	36.8
$Co_2D_{Y_7}Si_6O_{23}(OH)_{3}$	590	$\overline{2}$	20	cerite & apatite	
$Fe2La7Si6O23(OH)3$	700	$\overline{2}$	20	10.82	38.45
$Fe2Nd7Si6O23(OH)3$	550	$\overline{2}$	20	cerite <sup>a</sup>	
$Fe2Sm7Si6O23(OH)3$	550	$\overline{2}$	20	cerite <sup>a</sup>	
$Fe2Gd7Si6O23(OH)3$	680	$\overline{2}$	48	cerite & other phase	
$Mn_2La_7Si_6O_{23}(OH)_3$	680	$\sqrt{2}$	20	10.88	38.5
$Mn_2Ce_7Si_6O_{23}(OH)_2$	650	$\overline{2}$	10	cerite <sup>a</sup>	
$Mn_2Nd_7Si_6f_{23}(OH)_{3}$	550	$\overline{2}$	20	cerite & apatite	
$Mn_2Sm_7Si_6f_{23}(OH)_3$	550	$\overline{2}$	20	cerite & apatite	
$Mn_2Gd_7Si_6^b$	550	$\frac{2}{2}$	20	apatite	
$Mn_2Dy_7Si_6$ <sup>b</sup>	550		20	apatite	
$Mn_2Ho_7Si_6$ <sup>b</sup>	550	$\overline{2}$	20	apatite	
$Mn_{1.5}Ca_{.5}La_{7}Si_{6}O_{23}(OH)_{3}.\dots$	550	$\overline{2}$	20	cerite & apatite	
	550	$\overline{2}$	20	apatite	
	550	$\overline{2}$	20	apatite	
	1120	in air	20	10.31	37.25
	1120	in air	20	10.36	37.24
	1120	in air	20	10.45	37.32
	1050	in air	20	10.32	37.2
	1050	in air	20	10.34	37.2

<sup>&</sup>lt;sup>1</sup> Figures in brackets indicate the literature references at the end of this paper.

ing to the space group R3c  $[6]$ . X-ray powder diffraction data of Mg-La, Mg-Ho, and Co-La cerites and Mg-whitlockite,  $Mg_2Ca_7P_6O_{24}$ , were indexed and are given in tables 2 and 3.



Space group R3c, Hexagonal

Cerite was not found in previous studies of rareearth oxides-silica systems. Divalent ions play an essential role in the formation of cerite. The compositional range of cerites and a part of that of apatites in terms of ionic radii of the divalent ions and the rare-earth ions has been determined and is given in figure 1. It appears likely that the divalent and rareearth ion sites are ordered, and that no appreciable partial replacement of rare-earth sites by divalent ions takes place. Unit-cell dimensions of the cerite change only when the size of either the divalent or the rare-earth ion is changed, but not when the ratio of the amounts of two given cations in the source material is changed. A calcium end member of cerite phase does not exist under the present conditions, although the chemical analyses of natural cerite usually show several percent of calcium. The Mn analogs are converted to Mn-apatite at temperatures above 650 °C under 2 kbar  $H_2O$  pressure. The tentative boundary for cerite apatite conversion at 550 °C under 2 kbar is drawn in figure 1. The reverse reaction has not been observed, possibly because the reaction is too slow; but it may be that the equilibrium phase field for apatite is larger than indicated.

All the analogs of cerite were converted into apatite or apatite plus other phases after dehydration in air at temperatures above 900 °C. It was quite unexpected that the smaller rare-earth ions such as Dy or Ho also form cerite, because cerites found in nature contain exclusively large size rare-earth ions such as: La, Ce, Pr, Nd, and Sm.

Magnesium replacement in whitlockite was limited at the composition,  $Ca_7Mg_2P_6O_{24}$ . Starting material containing more Mg than corresponds to this end member gave an inferior whitlockite x-ray powder pattern without change in the unit-cell dimensions. With further increase of Mg in the charge, a second phase  $Ca_3Mg_3(PO_4)_4$  appears. The whitlockite solidsolution series is demonstrated in figure 2. Unit-cell dimensions vary as a function of Mg content.

Only a partial solid-solution series exists between cerite and whitlockite despite their close isomorphism







FIGURE 1. Compositional stability range of cerites. Abscissa is the ionic radii for divalent cations given by Wells (1962) [9], ordinate is rare earth radii given by Templeton and Dauben (1954) [10]

and close structural relationship. Large fields of stability for phospho-silico-apatites and monazites are present between the two compounds under the conditions used in the present experiments. Limited solid solution may extend from one toward the other. The



FIGURE 2. Whitlockite and Mg-whitlockite solid-solution series. Open circles=cerite, closed circles=apatite, crosses=mixed unidentified phases.

probable limits of solid solution may be around  $CaMg_2Ln_6Si_5PO_{23}(OH)_{3}$  for cerite, and  $Ca_8LnP_5SiO_{24}$ for whitlockite, but there is evidence to indicate that the solid-solution limit is a function of the size of rare-earth ion.

Infrared absorption spectroscopic analysis by R. Wilkins at Harvard University, showed abundant hydroxyl ion (up to several percent) in the synthetic Mg-La cerite possibly in two crystallographic sites  $(3690 \text{ cm}^{-1} \text{ strong infrared line and } 3660 \text{ cm}^{-1} \text{ medium})$ line). Synthetic Ca and Ca-Mg whitlockite, however were shown to be anhydrous as originally expected.

Europium and terbium activation of Mg-Gd analogs of cerite show typical red and yellow visible emission under a mercury fluorescent lamp, but their luminescent intensities are much lower than those displayed by the silicate apatities  $[1]$ .

From the results here presented, the formula  $M_2^2$ <sup>+</sup>Ln<sub>7</sub>Si<sub>6</sub>O<sub>23</sub>(OH)<sub>3</sub> seems in best agreement with the chemical analysis, infrared analysis, and synthesis of cerites. This formula is essentially the same as that given by Gay [6] and is isomorphous with Mg whitlockite except for two additional hydroxyl ions. However, calcium and magnesium ions in cerite do

not seem to be distributed randomly within rare-earth sites as suggested by Gay.

In the absence to date of a complete x-ray structure determination on cerite, the experimental results could be given an alternative interpretation by the following formulas:

 $M_3^{2+}$ Ln<sub>7</sub>Si<sub>7</sub>O<sub>27</sub>OH  $\cdot$  H<sub>2</sub>O or  $M_2^{2+}$ Ln<sub>8</sub>Si<sub>7</sub>O<sub>28</sub>  $\cdot$  3H<sub>2</sub>O. They correspond to that suggested by Keppler [8] for whitlockite. The latter formula was also that given by Glass et aL [7], from their extensive study on natural cerite from Mountain Pass, California\_

Support by the Advanced Research Projects Agency is gratefully acknowledged.

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